

1 Weather model verification using Sodankylä mast 2 measurements

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8 9 **Abstract**

10 Sodankylä, in the heart of Arctic Research Centre of the Finnish Meteorological Institute
11 (FMI ARC) in northern Finland, is an ideal site for atmospheric and environmental research
12 in the boreal and sub-arctic zone. With temperatures ranging from -50 °C to +30 °C, it
13 provides a challenging testing ground for numerical weather forecasting (NWP) models as
14 well as weather forecasting in general. An extensive set of measurements has been carried out
15 in Sodankylä for more than 100 years. In 2000, a 48 meter high micrometeorological mast
16 was erected in the area. In this article, the use of Sodankylä mast measurements in NWP
17 model verification is described. Starting in 2000 with the NWP model HIRLAM and
18 Sodankylä measurements, the verification system has now been expanded to include
19 comparisons between 12 NWP models and seven measurement masts, distributed across
20 Europe. A case study, comparing forecasted and observed radiation fluxes, is also
21 presented. It was found that three different radiation schemes, applicable in NWP model
22 HARMONIE-AROME, produced ~~during cloudy days~~ somewhat different downwelling long-
23 wave radiation fluxes during cloudy days, which however did not change the overall cold bias
24 of the predicted screen-level temperature.

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1 Introduction

Nocturnal and winter-time surface temperature inversions still pose a difficult challenge to weather forecast models. Various atmosphere to surface coupling issues are also problematic in climate models, especially at Arctic latitudes. –For the model development, versatile measurements are essential. The Arctic Research Centre of the Finnish Meteorological Institute (FMI ARC, <http://fmiarc.fmi.fi/>), is well suited for this purpose. The FMI ARC consists of two main stations, the headquarters in Sodankylä (67.368°N, 26.633°E), and the Pallas clean air research station (67.967°N, 24.117°E), which both provide ideal locations for atmospheric and environmental research in the boreal and sub-arctic zone.

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FMI-ARC dates back to the mid-nineteenth century when, in 1858, The Societas Scientiarum Fennica founded the first weather station in Sodankylä. Continuous meteorological measurements were started in 1908 and have been continued to this day (Savunen et al., 2014). Being accessible from all parts of the world, FMI ARC is also an excellent base for studying various themes of global change in a northern context.

Today, an extensive set of measurements ranging from basic meteorological data to heat and carbon fluxes as well as ozone and arctic snow coverage measurements is being performed at FMI ARC. Sodankylä observatory provides also facilities for receiving and processing polar satellite images, and FMI has conducted systematic aurora observations in the Finnish Lapland since late 1950's. The FMI ARC research sites belong to the Lapland Biosphere-Atmosphere Facility (LAP-BIAT, <http://www.sgo.fi/lapbiat/>), an infrastructure project through which the EU can fund visiting research groups. It has also been a site for various measurement campaigns (e.g., NOPEX/WINTEX campaign in 1997, Halldin et al., 2001), as well as various EU projects and measurement networks, (e.g. like CEOP (Savunen et al., 2014, http://data.eol.ucar.edu/master_list/?project=CEOP/EOP-3/4), CarboEurope IP (<http://www.carboeurope.org/>), and ICOS (<https://www.icos-ri.eu/>)).

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In the weather model verification, the traditional way is to perform detailed studies of model analyses and forecasts by comparing them with measurements afterwards. Another way to provide insight into model behaviour is to compare measurements with forecasts parallel with model runs in near-real time. Although based partly on less accurate (unchecked) measurements, this approach nevertheless provides valuable information about model behaviour and, when monitored frequently, can also act as a kind of alarm bell, alerting model

1 developers when there are apparent problems with model forecasts. Data collected this way
2 can also be used in model performance studies (Atlaskin and Kangas, 2006). As added
3 benefit, it provides means to monitor measurements.

4
5
6 Starting from 2000, the measurements at FMI ARC have been used to verify weather model
7 forecasts in near-real time. The verification was started with NWP model HIRLAM (Undén et
8 al, 2002; Eerola, 2013) and Sodankylä measurements, but has later been extended to cover
9 several other NWP models and mast measurement stations. Presently, a total of 12 models
10 and seven measurement masts are included. The models represent the activities of HIRLAM
11 (<http://hirlam.org>) and ALADIN (<http://www.cnrm.meteo.fr/aladin/>) NWP consortia, as well
12 as those of ECMWF (European Centre for Medium-Range Weather Forecast,
13 <http://www.ecmwf.int/>). The masts are located across Europe and run by various European
14 institutions. The forecast-measurement comparison plots with statistical analyses are provided
15 on-line as a part of HIRLAM forecast runs.

16 The harmonized and quality checked datasets collected in Sodankylä are also available for
17 more detailed research and model development. From the point of view of research, the most
18 valuable feature of the Sodankylä site is that it offers the possibility to combine various
19 simultaneous measurements, including those from a micrometeorological mast and a radiation
20 tower, as well as from dedicated snow and soil observations, AWS and atmospheric
21 soundings (see e.g. Coustau et al., 2014). In ~~this~~ the present article, these datasets are utilized
22 in a study of radiation from HARMONIE-AROME forecast system (Seity et al., 2011) versus
23 measured radiation in Sodankylä.

24 The Sodankylä measurements are likewise important in the initialization of NWP models in
25 operational forecasting. Of the measurements performed in Sodankylä, balloon soundings
26 (temperature, humidity, wind components) and some SYNOP measurements (surface
27 pressure, screen-level temperature, snow depth) are assimilated in the upper air and surface
28 analysis of HIRLAM and HARMONIE-AROME models.

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1 Section 2 contains description of Sodankylä site and Sect. 3 of the mast verification system. A
2 comparative study on HARMONIE-AROME radiation schemes is presented in Sect. 4, and
3 conclusions in Sect. 5.

4

5 **2 Sodankylä measurements**

6 The terrain around FMI ARC Sodankylä observatory (67.368°N, 26.633°E, altitude 179 m
7 asl, <http://fmiarc.fmi.fi/>) is moderately undulating, with isolated fells reaching up to 500 m
8 altitude. The observatory is located on the eastern bank of the river Kitinen, seven kilometres
9 southeast of the Sodankylä town centre, and about 100 kilometres north of the Polar Circle
10 and Rovaniemi. The vegetation in Sodankylä area is typical for the northern boreal zone, with
11 coniferous forest (mostly managed) and large open mires dominating the landscape. The
12 climate is characterised by long and cold continental-type winters and relatively warm but
13 short summers. During 1981-2010, the average yearly medium screen-level temperature was
14 -0.4 °C, yearly precipitation 527 mm, and snow cover duration 200 days (from 26 October to
15 14 May). The absolute minimum screen-level temperature during the same period was -49.5
16 °C and ~~with~~ absolute maximum value at +30.0 °C .

17 Due to the warming effect of the Gulf Stream the area can be classified as continental
18 subarctic or boreal taiga, by Köppen classification climate region Dfc (continental subarctic or
19 boreal (taiga) climates). However, with regard to stratospheric meteorology, Sodankylä can be
20 classified as an arctic site, often lying beneath the middle or the edge of the stratospheric
21 polar vortex and in a zone displaying intermittent polar stratospheric ozone depletion
22 (Savunen et al., 2014).

23 Continuous meteorological measurements have been performed in Sodankylä since 1908.
24 Ground-station observations every three hours record information on weather conditions
25 prevailing at ground level. In addition to standard weather observations, the basic
26 observational duties at the Observatory include regular recordings of solar radiation, sunshine
27 and hydrological quantities. Radiosonde measurements are carried out twice a day. During the
28 NOPEX/WINTEX measurement campaign, In 2000, a micrometeorological mast (48 m) for
29 atmospheric boundary layer measurements was erected in the area and has since been
30 producing data.

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1 ~~Sodankylä has also been extensively utilized for measurements in various projects, e.g.~~
2 ~~NOPEX and WINTEX in 1997 (Halldin et al., 2001), and CEOP (Savunen et al., 2014,~~
3 ~~http://data.eol.ucar.edu/master_list/?project=CEOP/EOP_3/4). During NOPEX/WINTEX an~~
4 ~~additional mast (18 m) was temporarily erected and used (Batchvarova et al., 2001). An~~
5 aircraft campaign to measure boundary layer properties was ~~also performed during~~
6 ~~NOPEX/WINTEX (Kangas et al., 19982001), the results of which were then used in a studies~~
7 ~~on satellite-based reflectance measurements (Kangas et al., 2001) and on regional momentum~~
8 ~~and sensible heat fluxes (Batchvarova et al., 2001).~~

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9 Data from most of the measurements is collected into a central data base at <http://litdb.fmi.fi/>.
10 It contains data not only from Sodankylä but also from other FMI ARC measurement sites. In
11 the following, the measurements used in the mast verification are briefly described.

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12 2.1 Micrometeorological mast

13 In 2000, a 48 meter high micrometeorological mast was erected in the immediate vicinity of
14 the Sodankylä observatory (<http://litdb.fmi.fi/micrometeorologicalmast.php>), and has since
15 been producing data. The height of the mast was limited by the presence of a near-by airfield.
16 It is located in a sparse Scots pine forest on a sandy podzol. The average tree height in is 12
17 m, tree density 210000 trunks per km², tree age 60-160 years, and the projected leaf area 1.2
18 m² (<http://en.ilmatieteenlaitos.fi/GHG-measurement-sites>).

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19 The mast is extensively instrumented with temperature, wind, humidity, and radiation
20 measurements at various levels (Fig.1, Table 12). The instruments used include
21 ~~PT100HMP155 (Pentronic Vaisala) thermometers for temperature, HMP35/45D (Vaisala)~~
22 ~~and humidity sensors, as well as and WAAA25 / WMT700 (Vaisala) and Thies 2D (Thies~~
23 ~~Clima) anemometers for wind speed and direction. Downwelling and upwelling short wave~~
24 and long wave radiation components (CNR4, Kipp&Zonen), net radiation (Nr-Lite,
25 Kipp&Zonen) and photosynthetically active radiation (PAR, LI190SZ, Licor) are measured
26 ~~at near~~ the top of the tower (458 m). Heat and momentum fluxes are measured at the 23 meter
27 level by the eddy covariance method (see more detailed description below).

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28 Additional near-ground measurements including soil temperature and ~~soil~~ moisture profiles,
29 soil heat flux, snow depth, and below canopy PAR are performed in the vicinity of the mast
30 (<http://litdb.fmi.fi/micrometeorologicalmastfield.php>) .

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1 2.2 Heat and momentum fluxes

2 The in situ fluxes of sensible heat, latent heat and momentum are measured at the
3 micrometeorological mast by the micrometeorological eddy covariance (EC) method, which
4 provides direct measurements of the fluxes averaged on an ecosystem scale. In the EC
5 method, the vertical flux is obtained as the covariance of the high frequency (10 Hz)
6 observations of vertical wind speed and the variable in question (temperature, H₂O
7 concentration, or horizontal wind speed) (Baldocchi 2003).

8 The eddy covariance measurement system at Sodankylä includes a USA-1 (METEK GmbH,
9 Elmshorn, Germany) three-axis sonic anemometer/thermometer and a closed-path LI-7000
10 (Li-Cor., Inc., Lincoln, NE, USA) CO₂/H₂O gas analyser. The measurements are performed at
11 23 m, 5 to 10 m above the mean forest height. The EC fluxes are calculated as half-hourly
12 averages taking into account the appropriate corrections. The measurement systems and the
13 post-processing procedures are presented in more detail by ~~and~~ Thum et al. (2009) and Aurela
14 et al (2015). See also Table 32.

15 2.3 Solar radiation tower

16 In addition to the basic synoptic measurements, a set of additional measurements is performed
17 on a 18 m high solar radiation tower in the observatory area. It contains measurements of
18 main radiation components: short wave radiation (CM11, Kipp&Zonen), direct normal
19 radiance (NIP, Eppley), long wave radiation (CG4 Kipp&Zonen) and aerosol optical depth
20 (PFR-N32, PMOD/WRC) (<http://litdb.fmi.fi/radiationtower.php>).

21 For consistency, all radiation data used in the mast verification is obtained from the radiation
22 tower. The measurements instruments on the radiation tower are also easily reachable and
23 allow more frequent maintenance than those on the micrometeorological mast. They are
24 quality-controlled and e.g. snow on the instruments is removed if found to exist. All
25 instruments except that for the outgoing LW radiation are ventilated. No heating is applied as
26 that would interfere with the measurements.

27 2.4 Automatic Weather Station

28 The automatic weather station (AWS) providing the official main weather parameters from
29 Sodankylä ~~AWS~~ has been in use since February, 2008. All the instruments and sensors at the

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1 station are calibrated annually. The parameters include screen-level temperature (PT100,
2 Pentronic) and humidity (HMP, Vaisala), air pressure (PTB201A, Vaisala), visibility (FD12P,
3 Vaisala), and cloudiness (CT25K, Vaisala). Wind speed and gust (WAA25, Vaisala) and wind
4 direction (WAV15, Vaisala) at the height of 22 m, as well as snow depth (SR50, Campbell
5 Scientific) are also provided (http://litdb.fmi.fi/apache2-default/luo0015_data.php).

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8 3 The mast verification system

9 3.1 Near-real-time comparison

10 Since 2002, near-real-time comparisons of model forecasts and in situ measurements have
11 been performed as a part of HIRLAM weather forecast model operational runs at FMI.
12 Started with HIRLAM forecast and Sodankylä measurements, the comparison has expanded
13 to comprise a total of 12 models and seven masts from around Europe. An eighth mast in
14 Estonia is presently being introduced into the system (Table 24). In addition to the direct on-
15 line comparison, long-term comparison statistics are provided. [Table 3 lists the parameters](#)
16 [included in the comparison.](#)

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17 To enable rapid update of the comparison, the comparison plots are produced as a part of the
18 operational HIRLAM forecast cycle (currently four times a day after synoptic hours 00, 06,
19 12, and 18 UTC) using the latest available data.

20 The HIRLAM program web site (<http://hirlam.org>) is used as the data pool, into which the
21 data providers transfer their data in prescribed format and from where it is retrieved by the
22 plotting routines located at FMI. The plotting is performed with Gnuplot
23 (<http://www.gnuplot.info/>) scripts, produced and run by the data retrieving program based on
24 perl and unix scripts.

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25 The parameters that are currently plotted include temperature, wind speed, and humidity at
26 specified levels as well as various heat and radiation fluxes (Table 32). With the original aim
27 in mind, the temperature difference between two metres and at a higher level (usually the first
28 model level) is also included in the plots as a measure of the surface temperature inversion.

29 ~~For all masts and models, the full set of parameters is not available, in which case an~~

1 ~~appropriate subset is plotted.~~ A sample plot showing screen level (2m) temperature from
2 HIRLAM forecast as compared to Sodankylä mast measurement (at 3m)s is shown in Fig.-2.

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3
4 ~~The~~An interactive web page ~~that has been set up~~ for browsing the comparison results has been
5 set up. The page enables side-by-side comparison of different mast/model combinations. is
6 visualised in Fig. 3. There are two panes, on each of which the user can select the desired
7 mast/model combination. By scrolling down the page, comparison for different parameters
8 can be viewed.

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9 Not all model-mast-parameter combinations are possible, however, because parameters
10 measured at different masts vary and all mast locations are not covered by all model
11 integration areas. In these cases, ~~a special "No comparison available" plot~~an appropriate
12 subset of the plots is shown. ~~The web page also contains~~Information about the parameters as
13 well as brief descriptions of the masts and models is also included ~~in the comparison~~. The
14 page is available to all HIRLAM and ALADIN consortia participants and to data suppliers as
15 a part of the general HIRLAM forecast visualisation pages.

16 3.2 Statistical comparison

17 Seasonal statistics compiled for individual observatories, or mast sites, containing the models
18 available at each respective station ~~are~~also calculated in the mast comparison as well.
19 Seasonal summaries of the daily comparisons, including a variety of descriptive and
20 comparative statistics, are shown under a separate heading on the interactive web page.

21 Graphs include time series of observed and modelled variables and the departures of model
22 output from the observations. They provide a qualitative view of how the models are doing,
23 and how their performance has varied during the season, thus linking model performance to
24 the prevailing conditions. These graphs are also useful for identifying gaps in the data.

25 Graphs of average model biases and rms-errors as function of forecast lead time serve to
26 quantify the errors, while scatterplots, histograms and mean diurnal cycles help to interpret
27 the errors physically by linking the average errors to specific conditions or hours of the day.

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28 As an example, Fig. 3 shows as the plots of RMSE and bias of screen-level (3 metres in the
29 mast) temperature and upwelling longwave-LW radiation (LWULWUP, obtained from the

1 [18-metre radiation tower, see Table 3](#)) for the spring period (March-April-May) of 2014. The
2 plots include [data from](#) four models, HIRLAM (FMI), HARMONIE-AROME (FMI), IFS
3 ([ECMWF](#)) and Arpege (Météo France) and they show the first 24 hours of the 00 UTC
4 forecasts. One can see that for [the FMI operational HIRLAM](#) there is a clear overestimation
5 of both ~~ULWUP~~ and the screen-level temperature. Here ~~LWULWUP~~ represents the
6 [surface temperature over open land in the measurements and that of the whole forest-covered](#)
7 [50-km² gridbox](#) grid box in the model. For HARMONIE-AROME and Arpege, we have
8 slight underestimation of both of these parameters, especially at about midday. For IFS, the
9 [correspondence between these two parameters is not so clear.](#)

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12 **4 Comparison of HARMONIE-AROME radiation fluxes to Sodankylä** 13 **observations: a case study**

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14 Spectrally averaged shortwave and longwave radiation fluxes at the surface are predicted
15 output variables of the contemporary NWP models. They are directly comparable to the
16 observed radiation fluxes, which could thus be used for the validation of the forecast along
17 with the near-surface temperature and humidity, anemometer-level wind, cloudiness, and
18 other variables diagnosed from the NWP model output in the standard station verification. In
19 particular, comparison of the simulated and observed radiation fluxes can give useful insight
20 for the development of the cloud and radiation parametrizations in the NWP models. Both in
21 reality and in the models, the short-term variability of the surface radiation fluxes is mostly
22 related to the variations of cloud and aerosol particles in air. In Sodankylä, the influence of
23 aerosol in the atmospheric radiation transfer is minor. In this section, we will test different
24 atmospheric radiation parametrizations in an experimental version of the HARMONIE-
25 AROME forecast system, based on [the reference cycle 38h1.2,](#)
26 [http://hirlam.org/index.php/hirlam-programme-53/general-model-description/mesoscale-
harmonie](http://hirlam.org/index.php/hirlam-programme-53/general-model-description/mesoscale-
27 harmonie)), against the Sodankylä radiation tower measurements.

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1 4.1 Measurements and numerical experiments

2 For a model-observation comparison, six components of radiation fluxes measured in the 18-
3 metre high Sodankylä radiation tower are available (Table 32): shortwave downwards
4 (~~SWDSWDN~~ or global radiation) and upwards (reflected), direct normal solar irradiance
5 (DNI), diffuse short wave solar radiation, long wave radiation downwards (~~LWDLWDN~~) and
6 upwards (~~LWUP~~). In this study, we compared the observed ~~SWDSWDN~~ and ~~LWDLWDN~~ to
7 their model counterparts for time period 15 January – 15 May 2014. The available one-minute
8 flux measurements were averaged over three-hour periods and compared with the three-hour
9 average fluxes derived from the accumulated radiation fluxes of the +6h and +3h
10 HARMONIE-AROME forecasts, which were initiated every 6h (00, 06, 12, 18 UTC). In
11 addition, the screen-level temperature observations provided by the Sodankylä automatic
12 weather station (AWS), representing the middle of each three-hour period, were selected for
13 comparison with the forecasted screen-level temperature. Sodankylä daily average
14 precipitation observations were extracted from FMI climatological data base.

15 The default atmospheric radiation parametrization of AROME (Seity et al., 2011) is based on
16 the radiation transfer code in the Integrated Forecast System (IFS cycle 25R1, European
17 Centre for Medium-Range Weather Forecast implementation in 2002), see ECMWF (2012)
18 and Mascart and Bougeault (2011), denoted here as ~~IFSRADifsradia~~. An alternative radiation
19 scheme originates in ALADIN (Mašek et al., 2015), hereafter denoted as
20 ~~ACRANEB2acraneb2~~. The radiation scheme of HIRLAM, based on Savijärvi (1990), see
21 also Nielsen et al. (2014), hereafter denoted as ~~HLRADIAhlradia~~, was is-available for
22 experimentation. All three schemes were tested within the framework of AROME physical
23 parametrizations by running three series of ~~experiments using a dedicated version (harmonie-~~
24 ~~38h1.radiation) of HARMONIE-AROME-experiments~~ over a domain covering Finland. A
25 horizontal resolution of 2.5 km and 65 levels in vertical were used. Lateral boundary
26 conditions for the experiments were obtained from the ECMWF analyses. For the initial state
27 of each +27h forecast, the objective analysis of the surface variables was combined with the
28 atmospheric analysis extracted from the boundary ~~files.Thefiles~~. ~~The surface-related~~
29 ~~parametrizations in AROME are taken care by the externalized surface scheme SURFEX~~
30 ~~(Masson et al., 2013)~~.

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4.2 Model – observation comparison in early spring 2014

~~Most of the winter days before mid-March 2014 were cloudy in Sodankylä. Most observed and predicted clouds were essentially non-precipitating. The non-precipitating clouds predicted by HARMONIE-AROME consisted mainly of (supercooled) liquid droplets while the ice crystal content was small. Some amount of (precipitating) snow and graupel was practically always present in the simulated clouds and some liquid/ice condensate at the lowest model level was often predicted. This is due to a recent change in cloud microphysics treatment in the HARMONIE reference system (Karl-Ivar Ivarsson, personal communication, 2015). Most of the days during 15 January – 15 March 2014 were cloudy in Sodankylä. Most observed and predicted clouds were essentially non-precipitating. The non-precipitating clouds predicted by HARMONIE-AROME consisted mainly of (supercooled) liquid droplets while the ice crystal content was small. Some amount of (precipitating) snow and graupel was practically always present in the simulated clouds. This is due to a recent change in cloud microphysics treatment in the reference system (Karl Ivar Ivarsson, personal communication). A small amount of liquid/ice condensate at the lowest model level was often predicted.~~

Every month, there were several days when more than one mm of precipitation, corresponding roughly to one cm of snowfall, was observed and predicted, while the first significant rainfall appeared in the end of April. These precipitation events were predicted well by the model. Falling precipitation was observed during the periods when also HARMONIE suggested significant snow and graupel content in the clouds. This indicates that in the model, most particles classified as precipitating indeed reached the surface, in agreement with the observations. Typically, the simulated condensate content of the precipitating particles was two to three times the liquid droplet water content, which in turn was an order of magnitude larger than that of the ice water content. In our experiments, only the cloud liquid droplets and ice crystals, but not the precipitating particles, were allowed to influence the radiative transfer in the atmosphere. This deviated from the default HARMONIE (cycle 38h1.2) settings, according to which ~~deviated from the reference system where~~ a fraction of the snow and graupel particles is accounted for when determining the cloud optical properties.

Figure 44 shows time-series of the observed and forecasted (+24h) screen-level temperature, SWDSWDN and LWDLWDN as well as the difference between the observed and

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Commented [1]: Keväällä 2014 käytössä oli 38h1; onko väliä?

Commented [2]: Reply to Carl Fortelius (10.03.2016, 12:48): "..."
Kokeet on ajettu 38h1.2:n erikoisversiolla ja havainnot haettu erikseen, ei ole väliä eikä mitään tekemistä operatiivisen kanssa.

1 ~~forecasetd~~forecasted LWDLWDN in February 2014. An overall cold bias of the screen-level
2 temperature forecast by the model using any radiation scheme was detected as compared to
3 the AWS observations (Fig. 44a). Typically, the forecast was one-two degrees colder than
4 observed.

5 In February, solar radiation flux (Fig. 44b) is small, Sodankylä being located north from the
6 polar circle. In February 2014, the maximum observed SWDSWDN value was ca 160 Wm⁻²,
7 while a typical daily maximum value was less than 80 Wm⁻². As the long-wave effects (Fig.
8 44c) are expected to dominate in the surface radiation balance, we will focus ~~on~~ the
9 LWDLWDN comparison.

10 Generally, the LWDLWDN flux was predicted well (Fig. 44 c and d). The largest differences
11 between predicted and observed LWDLWDN were found 1-2, 7-8 and 19-21 February. The
12 results were best when using the IFSRADjfsradia and ACRANEB2acraneb2 schemes, while
13 more deviations were found for HLRADIAhhradia.

14 Automatic weather station observations (not shown) indicated that during February 2014,
15 only the afternoon and night after the 20th was cloudless in Sodankylä. In this truly clear sky
16 case (both observed and simulated) all schemes correctly produced small LWDLWDN fluxes
17 and low screen-level temperatures. When observed clouds were not caught by the model,
18 LWDLWDN fluxes were underestimated by all schemes. This was the case e.g. on 21
19 February. Downwelling long-wave radiation was overestimated by HLRADIAhhradia (Fig.
20 44c, 44d) when the simulated clouds were optically thick (due to the assumed large super-
21 cooled liquid water content, not shown), for example during 9-12 February. During some
22 periods (7-8 and 17-19 February), the cold bias of the screen-level temperature was most
23 evident for HLRADIAhhradia, which showed the most underestimated LWDLWDN values
24 these days. Also the integrated cloud liquid water content was then smaller in the experiment
25 with HLRADIAhhradia than it was with other schemes. This might indicate secondary effects
26 due to the cloud-radiation interactions in the model. However, more studies are needed to
27 estimate the significance of this difference and to understand the mechanism behind it.

28 The simulated LWULWUP (Figure 44e) followed observations generally much more closely
29 than the screen-level temperature. This indicates that the surface (skin) temperature seen by
30 the radiation parametrizations was predicted well in most cases (with the exception of the first
31 two days and 7–8 February). In the model, the properties of the snow cover on ground and, to

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1 [some extent, the soil and vegetation properties under the snow, influence the surface](#)
2 [temperature and the grid-average LWULWUP.](#)

3 The different LWDLWDN produced by the different radiation schemes does not, however,
4 explain the systematic bias of the predicted screen-level temperature. LWDLWDN is a part of
5 the surface energy balance, which determines the (snow, soil) surface temperature that
6 interacts with the atmosphere. In the model, the diagnostic screen-level temperature is
7 obtained by interpolating between the predicted lowest model level (representing the layer up
8 to ca 28 metres from the surface) and the surface temperatures. In the interpolation, the
9 surface layer stability is taken into account. [The diagnostic estimation of the screen-level](#)
10 [temperature is likely to add uncertainty to the model-observation comparison. Thus, the](#)
11 [simulated screen-level temperature was evidently strongly influenced by the lowest model](#)
12 [level temperature, which in turn was dominated by the temperature advection in the low](#)
13 [troposphere.](#)

14 [In a model-observation comparison at a single location, phase errors of the large-scale](#)
15 [forecast in time and space show up if e.g. the arrival of an atmospheric frontal system has](#)
16 [been forecasted incorrectly. However, a systematic bias is hardly explained by the phase](#)
17 [errors. A comparison between the predicted lowest model level temperature with the](#)
18 [corresponding measurements of the micrometeorological mast, as well as a comparison](#)
19 [between the predicted surface temperature and the corresponding snow/soil surface](#)
20 [temperatures, might shed light on the problem. Predicted solar radiation fluxes, although](#)
21 [small in this period, deserve evaluation against the observations. This falls, however, outside](#)
22 [the scope of the present study.](#)

23 ~~The simulated upwelling long wave radiation (not shown), which represents the surface~~
24 ~~44) temperature, followed observations generally much more closely than the screen level~~
25 ~~temperature. This indicates that the surface (skin) temperature seen by the radiation~~
26 ~~parametrizations was predicted well in most cases (with the exception of the first two days~~
27 ~~and 7-8 February).~~
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1 ~~Thus, the simulated screen level temperature was evidently strongly influenced by the lowest~~
2 ~~model level temperature, which in turn was dominated by the temperature advection in the~~
3 ~~low troposphere. In a model observation comparison at a single location, phase errors of the~~
4 ~~large scale forecast in time and space show up if e.g. the arrival of an atmospheric frontal~~
5 ~~system has been forecasted incorrectly. However, a systematic bias is hardly explained by the~~
6 ~~phase errors. A comparison between the predicted lowest model level temperature with the~~
7 ~~corresponding measurements of the micrometeorological mast, as well as a comparison~~
8 ~~between the predicted surface temperature and the corresponding snow/soil surface~~
9 ~~temperatures, might shed light to the problem. Predicted solar radiation fluxes, although small~~
10 ~~in this period, deserve evaluation against the observations. This falls, however, outside the~~
11 ~~scope of the present study.~~

13 5 Conclusions

14 The near-real time mast verification of NWP forecasts, started in 2000, has proved to be very
15 useful in NWP model verification and, after being started with only one model and one mast
16 (HIRLAM and Sodankylä), has now expanded to include 12 forecasts and seven masts [across](#)
17 [Europe](#).

18 The mast verification system has been integrated with the [operational](#) runs of NWP model
19 HIRLAM, with data for other models and masts obtained through a common data pool. The
20 results are shown as a part of HIRLAM web-based visualisation pages that are available to all
21 data suppliers and members of HIRLAM and ALADIN NWP model consortia. The system is
22 not dependent on HIRLAM runs, though, and could be also run separately.

23 Statistics of the comparisons with e.g. long-term bias are also included in the verification,
24 although they are not updated daily but on seasonal basis. They provide seasonal summaries
25 of the daily comparisons, including a variety of descriptive and comparative statistics.

26 A comparative study of different radiation schemes applicable within HARMONIE-AROME
27 NWP system was also presented for [early](#) spring 2014. Based on this example, we conclude
28 that the three different radiation schemes produced generally good but somewhat different
29 [LWDLWDN](#) fluxes in cloudy days - and in February 2014, there was only one afternoon and
30 night free of clouds in Sodankylä. The [HLRADIA](#) [hrradia](#) scheme behaved most differently

1 from the other two schemes – [IFSRADIA](#) and [ACRANE2](#).
2 [HLRADIA](#) tended to overestimate [LWDLWDN](#) in case of optically thick clouds and
3 possibly underestimate it in case of optically thin clouds. However, when comparing the
4 simulated screen-level temperatures to those observed by AWS, the usage of any scheme
5 seemed to lead to a systematic cold bias of the order of one to two degrees. The reason of this
6 bias seems to lay outside the radiation parametrizations and –requires further study to be
7 understood.

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1 **Table 1.** Sodankylä micrometeorological mast measurements (see also Fig.1)

<u>Parameter</u>	<u>Measurement heights (metres)</u>
<u>Temperature</u>	<u>3, 8, 18, 32, 48</u>
<u>Humidity</u>	<u>3, 8, 18, 32, 48</u>
<u>Wind speed</u>	<u>18, 32, 38, 48</u>
<u>Wind direction</u>	<u>48</u>
<u>Global and reflected solar radiation</u>	<u>45</u>
<u>Long wave radiation up and down</u>	<u>45</u>
<u>Net radiation</u>	<u>45</u>
<u>Photosynthetically active radiation (PAR)¹</u>	<u>45</u>
<u>Snow depth</u>	<u>Ground level field</u>
<u>Precipitation</u>	<u>Ground level field</u>

2 ¹⁾ spectral range 400-700 nm

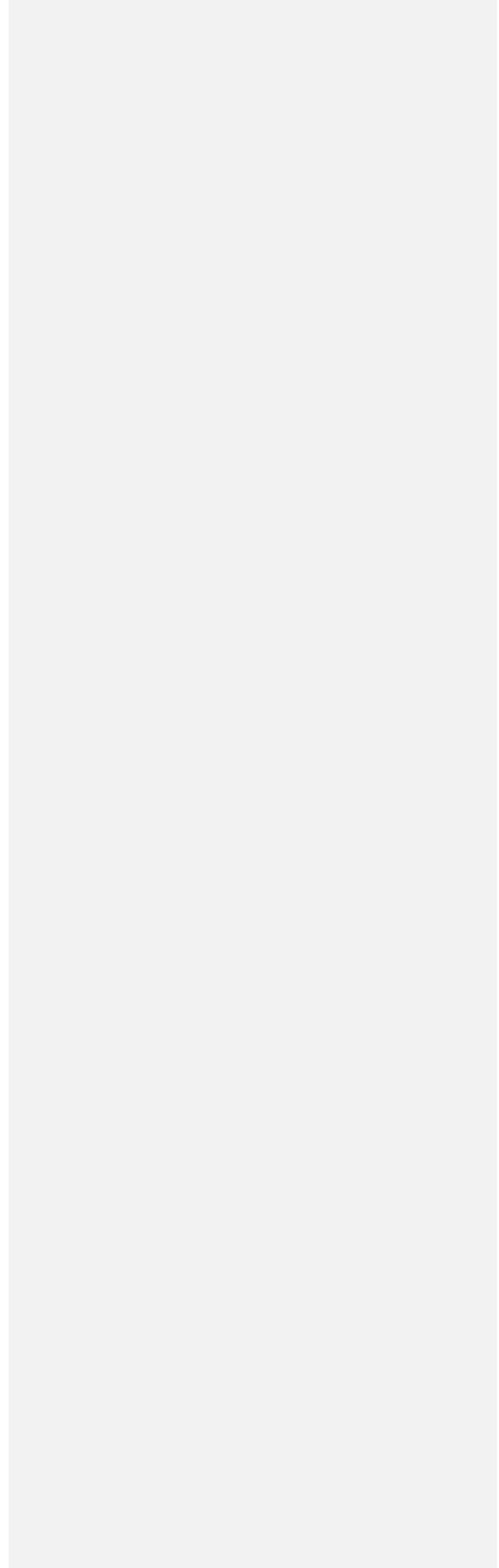
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2 **Figure 1.** Sodankylä micrometeorological mast (November 2015). T = temperature, WS =
3 wind speed, RH = relative humidity, T = temperature, WS/WD = wind speed/direction, SR =
4 solar radiation, GLOB = global radiation, REFL = reflected radiation, LWIN/LWOUT =
5 incoming/outgoing longwave radiation, SD = snow depth (A. Poikonen, 20156). See also
6 Table 1.

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9 **Figure 2.** Example mast verification plot from September 22, 2015: Screen-level (2m)2m
10 temperature from HIRLAM forecasts compared to Sodankylä mast measurement (3m
11 height)s. Red continuous line (OBS) shows measurements, dotted coloured lines (FCST)
12 show the first 24 hours from a set of consecutive forecasts.

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15 **Figure 3.** Statistical comparison of screen-level (3m in the mast) temperature and the
16 upwelling LW radiation for the first 24 hours of 00 UTC forecasts. Time period is March-
17 April-May, 2014, and the models HIRLAM (FMI), HARMONIE-AROME (FMI), IFS
18 (ECMWF) and Arpege (Météo France).

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20 **Figure 3.** Web page sample.

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23 **Figure 44.** Variables as function of time (x-axis, dates in February 2014 shown at the axis):
24 (a) screen-level temperature; (a) unit: °C; (b) SWDN; (cb) and LWDN-(e); (d), unit Wm⁻²;
25 difference predicted – observed LWDN; (ed) LWUP, unit Wm⁻². Temperature unit: °C, all
26 other in units of radiation fluxes (Wm⁻²). Colours of the curves and dots denote the observed
27 (red), ACRANE2_{acraneb2} (green), HLRAD_{hlradia} (grey), and HFSRAD_{ifsradia} (blue).

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The stated agreement between simulations and observations of upwelling long wave radiation could be shown. As the observations will include contributions from both the snow surface and trees, are they strictly comparable? It is also stated that comparison of the lowest model level temperature with mast measurements could shed light on the temperature bias problem; these measurements are available, so why not make the comparison?

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We added LWU as Fig.4e and modified the related text:

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“The simulated upwelling long wave radiation (LWU, Figure 4e)), which represents the surface temperature, followed observations generally much more closely than the screen level temperature. This indicates that the surface (skin) temperature seen by the radiation parametrizations was predicted well in most cases (with the exception of the first two days and 7-8 February). In the model, the properties of the snow cover on ground and, to some extent, the soil and vegetation properties under the snow, influence the surface temperature and the grid average LWU.”

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1 ~~We also added a reference to the SURFEX surface parametrizations to the end of 4. The~~
2 ~~suggested comparison between mast observations and the model's lowest level temperature~~
3 ~~falls out of the scope of the present study, which focuses on radiation fluxes. In fact this~~
4 ~~comparison would require significant additional data processing, both from the observations~~
5 ~~and from HARMONIE experiments.~~

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