



1 **"Shallow Geophysical Techniques to Investigate the Groundwater Table**
2 **at the Giza Pyramids Area, Giza, Egypt."**
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7 **ABSTRACT**

8 Geophysical studies were performed along selected locations across the Pyramids Plateau
9 to investigate the groundwater table and the near aquifer, which harmfully affected the existed
10 monuments of the Giza Pyramids and Sphinx. Electrical Resistivity Imaging (ERI), Shallow
11 Seismic Refraction (SSR) and Ground Penetrating Radar (GPR) techniques were carried out
12 along selected profiles in the plateau. Ten ERI, twenty six SSR and nineteen GPR profiles were
13 performed at the sites. The ERI survey shows that, the groundwater table is at elevations varying
14 from 13 to 18 m above the sea level (asl) and low resistivity values near the surface at the Great
15 Sphinx. ERI profiles, which were applied southeast of the Middle Pyramid (Khafre), show high
16 resistivity values near the surface, and water table is located at elevations ranging from 22 to 40
17 m asl, while the ERI profiles conducted south of Menkaure, show almost high resistivity near the
18 surface. The groundwater table is located at elevations ranging between 45 and 58m asl. The
19 aquifer layer shows electrical resistivities ranging between 10 and 50 Ohm.m. The considerable
20 high change in the groundwater table is due to the rapid increases of topography from the Great
21 Sphinx towards the Small Pyramids (Menkaure), where this part looks-like a scarp. The SSR
22 Survey is transmitted to know the different velocities and types of the layers, which can help in
23 knowing the saturated layers in the area. The GPR Survey is performed to delineate the water
24 table, which gives good matching with the ERI results.

25 *Keywords: Groundwater, Electrical Resistivity, Seismic refraction, GPR.*

26 **I. INDRDUCTION**

27 In recent years, the Egyptian Great Pyramids of Giza and Great Sphinx area are suffering
28 from the rising of groundwater table, due to the rate of water leakage from the surrounded areas,
29 due to the increase of expansion in urban developments of the new cities, population blocks,
30 gardens and agricultural expansion, as well as the water distribution from the nearby Nile River
31 Canals, such as El-Mansoria Canal. This problem promoted the need to monitor the groundwater
32 table in the Pyramids area, thus by using geophysical techniques we can evaluate the effect of
33 groundwater on the Pyramids and Sphinx area. Shallow geophysical investigation techniques for
34 site characterization and near-surface targets, with the integration of the available geological data
35 were used to develop a 3D model for the hydrologic system at the Giza Pyramids area. The
36 analysis and interpretation of this model lead to characterize the groundwater table in the



37 pyramids area, to define the possible sources of water leakage, and to evaluate the effect of the
38 groundwater on the pyramids and Sphinx area.

39 Geophysical studies play an important and effective role in groundwater investigation.
40 During the last few years, near-surface geophysics, especially Geoelectric Resistivity, Shallow
41 Seismic Refraction and Ground Penetrating Radar techniques have been widely used in Egypt to
42 characterize the groundwater table and the subsurface rock masses. The present work
43 demonstrates the integration of electrical resistivity Imaging (ERI), shallow seismic refraction
44 (SSR), and Ground Penetrating Radar (GPR) techniques to delineate the groundwater table in the
45 Giza Pyramids area, (Fig.1). The Giza Pyramids Plateau is composed of a limestone cliff,
46 changes abruptly from the other side to a sandy desert plateau. The Ancient Egyptians called this
47 place Imentet, "The West" or Kherneret, "the Necropolis". The three Giza Pyramids named
48 (Khufu), Chephren (Khafre), and Menkaure are located along this limestone Plateau.

49 **II. GEOLOGY OF THE AREA**

50 The Pyramids Plateau is formed from massive limestones and dolomites (nummulitic
51 wacke-packstones) of the Middle Eocene Mokattam Formation, which dips with about 5-10° to
52 the SE direction. Steep escarpments border the plateau to the north and east directions (Fig. 2).
53 Southwards, the Mokattam Formation is overlain by less resistant sandy marls, marls and
54 weakly cemented limestones (argillaceous mud-wackestones) of the Upper Eocene Maadi
55 Formation. The top unit of the Maadi Formation comprises several meters of massive, partly
56 dolomitized limestones (pack-grainstones) of the so-called "Ain Musa Bed". The Maadi
57 Formation shows a more gentle escarpment toward the Mokattam Formation, to the north and to
58 the eastern Nile valley alluvium. The present escarpments represent a Pliocene shoreline and
59 documented the transgression of the Early Pliocene Sea from the Mediterranean up to the pre-
60 Nile valley ("Eonile", Said 1981, 1982), after the largely continental Oligocene and Miocene
61 times (Blankenborn 1921 and Said 1962, 1981 and 1990). A thin wedge of the Pliocene
62 sediments rests discordantly on the Maadi Formation, but only veneer remains distinct against
63 the Mokattam escarpment (Fig. 2). The inferred fault along the Mokattam Formation of the
64 Pyramids plateau reflects the fracturing of the limestone.

65 **III. GEOPHYSICAL INVESTIGATION TECHNIQUES**

66 **III.1 Electrical Resistivity Imaging (ERI) Surveying and data acquisition**

67 Two-dimensional electrical resistivity imaging (tomography) surveys are usually carried
68 out, using a large number of electrodes, 24 or more, connected to a multi-core cable (Griffiths
69 and Barker 1993). Syscal-Pro resistivity meter, IRIS Instruments, France production, was



70 deployed at the site of the Giza Pyramids plateau using, the Dipole-dipole electrode array
71 configurations using 24 metal electrodes, with electrode spacing of 5 meter and 120 meter cable.
72 The length of spread is 115 m for each profile and maximum investigation depth is 23.5 m. The
73 automatic sequence were designed for the Dipole-dipole electrodes array configurations, using
74 Electre-Pro program, version V.2.02.0 of IRIS instruments, then uploading this sequence from
75 the PC to the Syscal-Pro resistivity meter, using the USB Dongle cable. In the field, select the
76 uploaded automatic sequence and start on the acquisition.

77 In the study area of the Giza Pyramids, ten electrical resistivity imaging profiles were
78 performed to characterize the resistivity values of the area, and hence to locate the groundwater
79 table. Table 1 and figure 3 show the location of the electrical resistivity imaging profiles
80 conducted at the study area. The profiles started from the Great Sphinx, through the Middle
81 Pyramid (Khafre) and end at the Small Pyramid (Menkaure) from the southern part of the Giza
82 Pyramids area.

83 **III.1.2 Data processing and interpretation**

84 The acquired electrical resistivity imaging profiles were processed and interpreted, using
85 Prosys II program (version V.3.02.08) of IRIS Instruments and Rse2Dinv (Version 3.59)
86 program of Geotomo software, Malaysia origin. Prosys II program is used to damp the data of
87 the geoelectric resistivity imaging from the Syscal-Pro resistivity meter to the PC, using the USB
88 Dongle cable, and utilized to filter and exterminate bad and noisy data acquired in the field.
89 Res2Dinv program applies the least square inversions on the data exported from the Prosys II
90 program, where the resistivity is plotted on a logarithmic scale function of the depth of the
91 subsurface.

92 Ten electrical resistivity imaging (ERI) profiles were performed over the study area of
93 the Giza Pyramids plateau. The topographic elevation is considered for each ERI profile and fed
94 to the Res2Dinv program. The interpretations of the ERI1 to ERI3 profiles, which were taken
95 beside the Great Sphinx, shows low resistivity values near the surface and shallow water table,
96 which lies at elevations ranging from 13 to 18 m asl. The interpretation of the ERI4 and ERI5
97 profiles, which were located southeast of the Middle Pyramid (Khafre), shows high resistivity
98 values near the surface, where the water table is followed at elevations range from 15 to 43 m
99 (Figs. 4a and 4b). The analysis of the ERI8 to ERI10 profiles, which were conducted south of the
100 Small Pyramid (Menkaure), shows almost high resistivity near the surface. The water table is
101 located at elevations ranging between 45 and 58 m (Figs. 4c, 4d and 4e, respectively). These ERI
102 Models reveals mostly four layers in average in most parts of the study area.



103 **III.2 Shallow Seismic Refraction (SSR)**

104 Compressional waves or (P-waves) are used almost exclusively in the seismic exploration
105 for both seismic reflection and refraction. Especially at shallow depths, we primarily are
106 concerned with P-wave velocities in the rocks (consolidated materials) and sediments
107 (unconsolidated materials). In shallow refraction work, the P-wave velocities are often sufficient
108 to describe the ground layers in terms of dry and wet overburden, and fresh and weathered
109 bedrock. There are no unique velocity values for rocks or sediments; however, a few general
110 rules are suggested by these values (Burger 1992). The water saturation, porosity, weathering,
111 fracturing and compaction are factors affecting the layer velocities.

112 In the present study, however, the main target for applying the seismic refraction
113 technique is to determine the seismic velocities and thicknesses of the different successive layers
114 and to trace the lateral distribution of the subsurface layers throughout the investigated area.
115 Depth-velocity models are constructed; these models reflect the number of layers penetrated by
116 the seismic waves. Also, the type of lithology of each layer and water table within layer are
117 determined, according to the values of velocities of the seismic waves through layers.

118 ***III.2.1 Refraction data acquisition and survey parameters***

119 Twenty-six shallow seismic refraction profiles were acquired at the study area (Fig. 3).
120 OYO McSEIS-SX seismograph, of 24 geophones-channels, was deployed at the studied site to
121 collect the seismic refraction data. A sledge hammer (of 10 Kg) and an iron plate are used as P-
122 wave seismic source. The used inter-receiver distance is 5 m. The numbers of shots are 5 shots
123 per spread. Two off-set shots (each 20 meters from each end), forward (5 m from the first
124 geophone), reverse (5 m from the last geophone) and a split spread shot. The spread, performed
125 by seismograph covers 115 m. The nature of these important historical and touristic site in
126 Egypt, where a huge number of visitors and human activities existed in the site, is imposing a
127 considerable amount of noises to the recorded data. These noises were minimized as possible by
128 using the internal frequency domain filter applied by the seismic seismograph. To enhance the
129 data quality and decrease the random noises, each shot was repeated several times and stacked.

130 ***III.2.2 Refraction data interpretation***

131 The first arrivals of the collected P-waves are picked, using Pickwin of SeisImager
132 software version 4.2 (OYO 2011). These picked data are interpreted to get the depth-velocity
133 models, by applying appropriate inversion techniques. Figures (5, 6, 7 and 8), respectively show
134 examples of the interpreted seismic refraction profiles conducted at the site area. A three layers
135 model assumed to represent the subsurface succession with the inverted velocities and
136 thicknesses. The top most layer exhibits a velocity range of 400-900 m/s, and is correlated with



137 loose dry sand, fill and debris (which is corresponding to the first and second layer of electrical
138 resistivity model). The thickness of this layer ranges between 2 and 5 meters. The second layer
139 shows a velocity range between 1200 and 2400 m/s, this layer is correlated with wet and
140 saturated sand (which is corresponding to the third layer of electrical resistivity model). The
141 thickness of this layer varies from 10 to 20 meters. The third layer shows a higher domain of
142 velocity, where it ranges between 2800 and 3800 m/s, which can be correlated to marly
143 limestone and limestone (corresponding to the forth layer of electrical resistivity model), which
144 is considered as the aquifer layer, Table 1.

145 **III.3 Ground Penetrating Radar (GPR) techniques**

146 Ground-penetrating radar (GPR) is an effective tool to visualize the structure of the
147 shallow subsurface. Ground penetrating radar (GPR) has become a popular tool in the
148 environmental and engineering studies for the near-surface targets (Jol and Bristow 2003). It is a
149 non-invasive geophysical technique designed primarily for subsurface investigation (Neal 2004;
150 Comas et al. 2004). A GPR system detects changes in the electrical properties of the shallow
151 subsurface, using discrete pulses of high frequency electromagnetic (EM) energy, usually in the
152 10-1000 MHz range (Neal 2004).

153 The technique has been successfully applied in a wide range of environmental studies,
154 however an understanding of the capabilities and limitations of GPR is vital when considering
155 using the technique, with the quality of GPR results often being dependent on the surveyed
156 environment (Daniels 2004).

157 **III.3.1 GPR Surveying of the study area**

158 The GPR survey was carried out with MALA ProEx of Mala Geosciences, Sweden, using
159 100 MHz shielded antenna as a central frequency and data displayed, using a laptop computer.
160 Nineteen GPR profiles were performed along the study area of Giza Pyramids. The lengths of
161 GPR profiles range from 40 to 200 m, a total of about 2.5 kilometer of GPR surveys were
162 operated at the site. Locations and directions of the GPR profiles are viewed in Figure 3 and
163 Table-1. Surface topography elevations were taken into account in the GPR surveys, which can
164 be corrected in the radar processing, using static corrections. Wheel calibration was made near
165 the Great Sphinx along 30 m in distance, the velocity used in calibration is 100 m/ μ s and the No.
166 of stacking equal to 16. The depths of penetration vary from 8.5 to 20 m.

167 **III.3.2 GPR data Processing and analysis**

168 GPR data are subjected to a scheme of signal data processing, using Reflex-Win Version
169 6.0.9 software to enhance the quality of the gained data. The GPR data are displayed in cross
170 sections with the distance along the profile for the X-axis and with the two-way travel times of



171 the reflected GPR waves for the Y-axis. To convert the time sections into depth sections, an
172 average velocity of 0.1 m/ns was used, assuming a possible variation in depth of +/-10%., the
173 ground-vision of Mala Package and Reflex-W package are furnished to facilitate the processing
174 and interpretation of the acquired GPR data.

175 GPR data processing corrects the start time to compensate for air-wave and contact with
176 the ground, a DC-shift filter and an amplitude correction (Dewow) were applied to remove the
177 constant offset and compensate the loss spread and attenuation, respectively. Static corrections
178 were applied to the data to compensate for changes in the topographic elevations. A band-pass
179 filter, 2-D running average and gains function were applied to enhance the amplitudes of signals.
180 The background removal filter was applied to the data, the filter performs a subtraction of an
181 averaged trace. Deconvolution and stacking were performed to enhance the signal to noise ratio,
182 while Kirchhoff and diffraction migration processes were applied to the data to correct the
183 positions of reflection points. Muting also was introduced in some radar sections to remove the
184 bad data.

185 ***III.3.3 Interpretation of GPR Data***

186 The different colors of the radargram reflect the amplitudes of the reflected EM waves,
187 which are an indication of the change in the subsurface layers conductivities and dielectric
188 constants. GPR data resolve the locations of the layers boundaries as the dielectric constants of
189 the compositions changed, they are delineating the depths and extensions of the layers. GPR also
190 mapped the water table at the site. Nineteen GPR Profiles were conducted in the study area of the
191 Giza Pyramids (Fig. 3). In a way of knowing the groundwater table, the interpretation of GPR
192 profiles subdivided the area into four parts, according to the nearest one in distance and elevation
193 topography.

194 Area-I comprises the GPR Profile-1, which is located to the northwest of the Great
195 Pyramid. The site is a low land followed a scarp, it looks-like a wadi behind the plateau. The
196 interpreted GPR profile-1 reveals that, the water table uncertainly might be located at an
197 elevation of 20 to 21 meters asl, where the ground surface at 30 to 31 meters asl. Area-II
198 involves the GPR profile-2 to GPR profile-5 and the GPR profiles 9 and 10. Such GPR profiles
199 are conducted along the low land of Nazlet El-saman village, beside the wall and east of the
200 Great Pyramid and Sphinx. The interpreted profiles, despite the noisy nature of the site due to the
201 walls and houses, reveals that the water table might be located at elevations ranging from 11.5
202 to 17.5 meters asl. Area-III includes the GPR Profiles 6 to 8, which are located along the southern
203 eastern part of the study area, near El-Gabal El-Kebly. The water table is interpreted to locate at
204 elevation about 13 to 13.5 meters asl. Area-IV comprises the GPR profiles 11 to 19, which are



205 located at the southern part of both the Middle Pyramid (Khafre) and the Smallest Pyramid
 206 (Menkaure). The water table is located at elevations ranging from 14.5 to 68 meters asl. Table-1
 207 summarizes the results of water table (WT) elevations to all the GPR profiles of the study area.
 208 Figures (9a to 9d) view examples of the interpreted GPR profiles for each part.

209 *Table-1: summarized average ground elevations and the interpreted Groundwater table elevations with Coordinates for the*
 210 *different geophysical measurements*

Geophysical survey			Coordinates		Direction	Average Ground Elevation (m)	Average Water Table Elevation (m)
ERI No.	Seismic No.	GPR No.	X	Y			
	SSR1	GPR1	320446	3317418	NE-SW	30	20
	SSR2		320356	3317386	NE-SW	31	21
	SSR3	GPR2	320228	3317226	N-S	26	17
	SSR4		320110	3317236	N-S	27	16
	SSR5	GPR3	320028	3317179	NW-SE	20	12.5
	SSR6		319845	3317010	NW-SE	20	11.5
	SSR7	GPR4	319739	3316914	NW-SE	21	13
	SSR8	GPR5	319582	3316803	NW-SE	18	11.5
	SSR9	GPR6	319392	3316898	E-W	17	13.5
	SSR10	GPR7	319238	3316958	E-W	17	13
	SSR11		319403	3318602	E-W	18	13.5
	SSR12	GPR8	319344	3318544	E-W	18	13
ERI 1	SSR13	GPR9	320463	3317802	N-S	21	15-15.5
ERI 2	SSR14	GPR10	320455	3317716	NW-SE	20	15-15.5
ERI 3	SSR15	GPR11	320581	3317211	NE-SW	20	14-15
ERI 4	SSR16	GPR12	320542	3317251	NW-SE	27	21-22
ERI 5	SSR17	GPR13	320462	3317336	NE-SW	35	27-28
ERI 6	SSR18	GPR14	320394	3317358	N-S	42	37
ERI 7	SSR19	GPR15	320753	3316848	NE-SW	50	42-43
ERI 8	SSR20	GPR16	320713	3316845	N-S	58	47-48
ERI 9	SSR21	GPR17	320667	3316846	NW-SE	64	57
ERI10	SSR22	GPR18	320625	3316845	NW-SE	66	59
	SSR23	GPR19	320441	3317414	NW-SE	75	68
	SSR24		320370	3317381	NW-SE	81	70
	SSR25		320284	3317316	NE-SW	99	90
	SSR26		320169	3317128	NE-SW	102	93

211 IV. INTEGRATION OF THE DIFFERENT GEOPHYSICAL TECHNIQUES

212 IV.1 Comparison among the geophysical results at Area-I

213 The correlation among the SSR and GPR data at Area-I of the study area of Giza
 214 Pyramids plateau (Fig. 3) shows that, the results obtained from SSR and GPR data are were quite
 215 matched, where the interpreted water table (WT) from the SSR is located at elevation of 21 m,
 216 while the water table interpreted from the GPR is located at elevation of 20 m.



217 ***IV.2 Comparison among the geophysical results at Area-II***

218 The correlation between ERI, SSR and GPR data from some selected profiles at Area-II
219 of the study area of Giza Pyramids plateau (Fig. 3) views relatively the same results, where the
220 interpreted water table (WT) from the ERI is at elevation 15.5 m, while the water table
221 interpreted from the SSR is located at elevation 15 m, and the interpreted water table from the
222 GPR locates at elevation 15 m.

223 ***IV.3 Comparison among the geophysical results at Area-III***

224 The correlation among the SSR and GPR data at Area-III of the study area of Giza
225 Pyramids plateau (Fig. 3), exhibits good matching among the different techniques, where the
226 interpreted water table (WT) from the SSR is at elevation 13.5 m, while the water table
227 interpreted from the GPR is located at elevation 13 m.

228 ***IV.4 Comparison among the geophysical results at Area-IV***

229 The correlation among the ERI, SSR and GPR data from some selected profiles at Area-
230 IV of the study area of Giza Pyramids plateau (Fig. 3) gives results were quite matched, where
231 the interpreted water table (WT) from the ERI is at elevation 22 m, while the water table
232 interpreted from the SSR is located at elevation 22 m, and the interpreted water table from the
233 GPR locate at elevation 22 m. Figure 10 views an example of the correlations among the ERI,
234 SSR and GPR data from some selected profiles at Area-IV of the study area of Giza Pyramids
235 plateau.

236 ***IV.5 2-D and 3-D presentation of the geophysical interpretation***

237 From all the geophysical techniques (ERI, SSR and GPR) that applied at the study area of
238 Giza Pyramids, the groundwater table elevations are interpreted and shown in Table-1 and Figure
239 11a. Figure 11a represents the groundwater elevations map from the geophysical surveys, which
240 were applied at the study area of Giza Pyramids in 2016, posting on it the groundwater levels
241 from the borehole Piezometers, which were installed by Cairo University in 2008 prior to the
242 activation of the dewatering system by AECOM 2010. All the previous geophysical surveys
243 conducted in 2016 show that, the pumping system by AECOM 2010 has good effect in lowering
244 the groundwater levels form the piezometer No.1 to the piezometer No.11, but from piezometer
245 No.12 to piezometer No.16, which represent the area concentrated around the Great Sphinx, still
246 have high groundwater levels and need more withdrawal to the groundwater in this part. Also we
247 recommended, making ceil or barrier in this part because of the surface is saturated with water,
248 which may be due to lake of the impervious clay in the subsurface in this region. Figure 11b
249 integrates the different geophysical techniques of the ERI, SSR and GPR into a 3D Model. The
250 3D model illustrates the topography, number of layers and the interpreted groundwater table
251 elevations above the sea level of the study area of the Giza Pyramids plateau.



252 **V. Conclusions**

253 The interpretation of the electrical resistivity imaging (ERI) survey, the aquifer layer
254 shows electrical resistivities ranging between 10 and 50 Ohm.m. The imaging profiles near the
255 Great Sphinx show the groundwater table at elevations varying from 13 to 18 m asl. The imaging
256 profiles applied southeast of the Middle Pyramids (Khafre) show high resistivity values near the
257 surface, and the groundwater table is located at elevations range from 22 to 40 m asl. The
258 imaging profiles conducted to the south of the Small Pyramids (Menkaure) reveal almost high
259 resistivity near the surface, where the groundwater table is located at elevations varying between
260 45 and 58 m asl.

261 Twenty six shallow seismic refraction (SSR) spreads were conducted across the study
262 area of Giza Pyramids plateau, with spreads are ranging in length between 95 and 155 meters. A
263 model of three layers assumed for the shallow section of the study area. The top most layer
264 exhibits a velocity range of 400-900 m/s, this layer is correlated with loose dry sand and fill, with
265 a thickness ranges between 2 and 6 meters. The second layer shows a velocity range between
266 1200 and 2800 m/s, this layer is correlated with wet and saturated sand, with a thickness varies
267 from 5 to 12 meters, where the groundwater level is raised from the deep aquifer of the limestone
268 up in the second layer. The third layer shows a wide variation in velocity, where it ranges
269 between 2400 and 3950 m/s, which can be correlated with marly limestone and limestone.

270 Nineteen GPR profiles were performed along the study area of Giza Pyramids plateau,
271 with a total length of about 2.5 kilometer of GPR surveys. In Area-I, the groundwater table is
272 interpreted at elevations ranging between 20 and 21 m. In Area-II, the groundwater table is
273 interpreted at elevations varying from 11.5 to 17.5 m. In Area-III, the groundwater table is
274 interpreted at elevations ranging between 13 to 13.5 m. In Area-IV, the groundwater table is
275 interpreted at elevations varying from 14.5 to 22.5 m around the Sphinx part and at the part of
276 high topography near the Pyramids of Khafre and Menkaure, where the groundwater table is
277 interpreted at elevations of 32 to 68 m asl.

278 The groundwater table is interpreted from the acquired geophysical data along the
279 conducted profiles to be at the following elevations: Area-I: The groundwater table is located at
280 elevations of 20 to 21 meter asl; Area-II: The groundwater table is located at elevations of 11.5
281 to 17.5 meters asl; Area-III: The groundwater table is located at elevations of 13 to 13.5 meters
282 asl and Area-IV: The groundwater table is located at elevations of 14 to 58 meters asl.

283 It is evident that, the water table rises up the marly limestone layer of the Plateau, which
284 is the formation of the Pyramids, that causes a serious problem, due to the corrosion and
285 dissolution of the monuments. It is recommended to lower the water table to a level below the
286 marly limestone to protect the existed monuments.

287 **References**



- 288 **AECOM, (2010):** Groundwater Modeling and Alternatives Evaluation. Pyramids Plateau
289 Groundwater Lowering Activity.
- 290 **Blankenborn M., (1921):** Aegypten, Handbuch der regionalen Geologie. Heidelberg: Carl
291 Winters Universitätsbuchhandlung.
- 292 **Burger H. R., (1992):** Exploration geophysics of the shallow subsurface, Prentice and Hall, 489
293 p.
- 294 **Cairo University Oct., (2008):** Operation of Production Wells and Piezometer Readings below
295 the Sphinx and Wadi Temple for the Period from 15/7/2008 to 15/10/2008.
- 296 **Comas X., Slater L. and Reeve A., (2004):** Geophysical evidence for peat basin morphology
297 and stratigraphic controls on vegetation observed in a northern peat land. Journal of
298 hydrology 295, pp:173-184.
- 299 **Daniels D.J., (2004):** Ground penetrating radar (2nd edition). The Institution of Electrical
300 Engineers: London.
- 301 **Dobrin M. B., (1976):** Introduction to geophysical prospecting. McGraw- Hill Book Co., New
302 York, USA, 630 p.
- 303 **Griffiths D. H. and King R. F., (1965):** Applied geophysics for Engineering and geologists,
304 Pergamon press, Oxford, New York, Toronto, 221p.
- 305 **Jol H. M. and Bristow C. S., (2003):** GPR in sediments: advice on data collection, basic
306 processing and interpretation, a good practice guide. In Ground penetrating radar
307 insediments, Bristow CS and Jol HM (eds). Geological Society: London, Special Publication
308 211; pp: 9- 28.
- 309 **Loke M. H., (2012):** Tutorial: 2-D and 3-D electrical imaging surveys. Course Notes.
- 310 **MALÅ ProEx, (2011):** Operating Manual v. 2.0. MALÅ Geoscience AB, Sweden.
- 311 **Neal A., (2004):** Ground-penetrating radar and its use in sedimentology: principles, problems
312 and progress. Earth science reviews 66: pp: 261-330.
- 313 **OYO, (2011):** Instructor manual of McSEIS-SX24 device. OYO Company, Japan.
- 314 **Parasins D. S., (1979):** Applied geophysics: 3rd ed., Capman and Hall, London, 269 P.
- 315 **Said R., (1990):** The Geology of Egypt (Rushdi Said ed.). Rotterdam: A.A. Balkema.
- 316 **Said R., (1981):** Geological evolution of River Nile. Springer-Verlag.
- 317 **Said R., (1962):** Geology of Egypt, New Amsterdam, Elsevier Pub. Co.
- 318 **Sharma P. V., (1997):** Environmental and engineering geophysics, United Kingdom, Cambridge
319 University Press, 475p.
- 320 **SYSCAL Pro, (2011):** User's manual of 10 channels Resistivity-meter for Resistivity and IP
321 measurements, IRIS Instrument, France.
- 322 **Yehia A., (1985):** Geological structures of the Giza pyramids plateau. Middle East Res. Center,
323 Ain Shams Univ., Egypt, Sci. Res. Series, 5, pp: 100-120.