Response to Referee #1

This manuscript investigated climate change effects on extreme temperatures in the Blue Nile Basin. After a description of dataset and case study, authors described applied methodologies, among which emerges the widely employed Mann-Kendall test. The investigation was carried out by applying these tools to extreme temperatures detected in Blue Nile Basin. The topic is of paramount importance for hydrological applications to be implemented in the areas covered by this study. However, to my opinion the paper needs substantially improvements, both in its structure and results analysis, that are of major importance.

response: thank you very much. Your substantial and very appreciated improvements were done in the revised version of the paper

2 General comments: The paper deals with a relevant topic for modern hydrology, highlighting the need of such analyses for a wide field of real applications. However, to my opinion the paper needs to be improved substantially in the introduction and in the result analysis, and the abstract reformulated in a more readable way. In particular, the introductive section should be rebuilt in order to provide a more logical discussion about the general framework and local situation, clearly specifying motivations and goals of the paper. More detailed comments will be shown in specific comments section. response: the abstract was reformulated (lines 10-22) and the motivations and goals of the paper were specified clearly in the revised version of the paper (lines 103-107) ***all line numbers correspond to the tracked changes version of the paper to facilitate revision**

The second – and, to me, more important – issue is related to the analysis on Mann-Kendall test results and arise from the citation of the paper of Yue et al. (2002) at line 168. This is one of the most famous papers (1048 citations on Scopus ad November 17, 2020) on the use of Mann-Kendall test for detecting trends in hydrological series, highlighting the role of power evaluation when applying this test. In this way, a complete analysis of applications to real data can be performed, investigating both type I and II errors. A lot of papers discussed on the practical implications that the assessment of power can generate. Among the most recent, I suggest to refer to the following for the specific reference to Mann-Kendall (published in 2020): - Totaro, V.; Gioia, A.; Iacobellis, V. Numerical investigation on the power of parametric and nonparametric tests for trend detection in annual maximum series. Hydrol. Earth Syst. Sci. 2020, 24, 473–488. - Wang, F.; Shao, W.; Yu, H.; Kan, G.; He, X.; Zhang, D.; Ren, M.; Wang, G. Reevaluation of the Power of the Mann-Kendall Test for Detecting Monotonic Trends in Hydrometeorological Time Series. Front. Earth Sci. 2020, 8.

This about an application of power evaluation with parametric Likelihood Ratio test: -Németh, L., Hübnerová, Z., Zempléni, A. Comparison of trend detection methods in GEV models. Communications in Statistics-Simulation and Computation, 2020, 1-16. And these latter to implications and concerns on the need of evaluating the power: - Vogel, R. M., Rosner, A., and Kirshen, P. H.: Brief Communication: Likelihood of societal preparedness for global change: trend detection, Nat. Hazards Earth Syst. Sci., 2013 13, 1773–1778. - Serinaldi, F., Kilsby, C. G., and Lombardo, F.: Untenable nonstationarity: An assessment of the fitness for purpose of trend tests in hydrology,. Clarified this issue, I know that lots of hydrological applications are carried out in the same way as you did. However, to my opinion you should base your findings reporting some notes supported by literature references about Mann-Kendall test power, to reinforce your statements, in order to provide a more complete and appropriate interpretation of results.

response: the suggested references were studied and the current paper was enhanced using these papers in the revised version of the paper (lines 199-212)

Finally, in addition to these two remarks, I would like to see a more detailed discussion in Par. 3, that I found too short and an inversion between order of Par. 2.1 and 2.2.. response: more detailed discussion in Par. 3 were done (lines 199-303) and an inversion between the order of Par. 2.1 and 2.2in was presented in the revised version of the paper (lines 134-178)

Specific comments: - Line 13: specify which data; to me, monthly is too generic and is repeated at lines 13-14; - Lines 15-17: please, provide a clearer summary of your methodologies;

response: the repetition of monthly in the abstract was revised and a clearer summary of the methodologies was provided in the revised version of the paper (lines 10-15)

- Line 28: there has been: : : I think that it still is; response: the sentence was updated in the revised version of the paper (line 30)

- Line 29: its direction. Please, use a more specific term; response: the word direction was replaced with (trend) (line 31)

- Line 30: and the potential river basins in the Nile Basin: what do you mean? response: the authors mean that this sub-basin has the potential for development and hydropower generation (lines 33-34)

- Line 35: please, remove comma; response: was removed (line 39)

- Lines 39-40: this statement should be moved to the Conclusion section; response: the statement will be moved to the Conclusion section (lines 365-367)

- Lines 42-44: please, support your statement with references; response: the reference (Collins et al., 2019) was used (line 49)

- Line 46: what consequences are you referring to? – response: the consequences as higher evaporation rates and deterioration of quality

Lines 48-51: please, support your statements with references;

response: the reference (Sohoulande Djebou and Singh, 2016) was used (line 57)

- Line 59: it is not clear what basins are you referring to; response: the basins studied in the papers of (Anache et al., 2018, Bergström et al., 2001, Chen et al., 2012, Yan et al., 2020)

- Lines 63-64: is this statement referred to the work of Gleick (2000)? response: yes, the statement referred to Gleick (2000) work

- Lines 71-72: to me, you can better specify the type of variable you are analyzing; response: the variables are min, mean, and max monthly temperature (line 79)

- Lines 79-86: this detailed discussion should be moved to case study description, leaving only some notes about climate of Ethiopia that are strictly essential for developing the introduction; response: this discussion was moved to the case study description section (lines 143-150)

- Line 103: please, remove &; response: was replaced by "and" (line 160)

- Lines 104-107: to me, you have to provide a better declaration of hydrological variables you are investigating and to which you are applying tests; response: hydrological variables were declared (lines 162)

- Lines 108-110: please, report more details on the occurrence and treatment of missing data; response: more details on the occurrence and treatment of missing data was introduced in the revised version of the paper (lines 166-168)

- Line 118: Figure 1, pleas improve readability of words and numbers (e.g., increase dimensions); response: the figure was updated in the revised version of the paper (line 177)

- Lines 142-144: please, provide a reference for your statement; response: the references (Onyutha, 2017, Partal and Kahya, 2006) were used (line 189)

- Lines 153-154: I think you can rephrase your statement in a clearer way. I can't understand what do you mean;

response: the statement will be rephrased to (Moreover, MK and Sen's slope estimator test were employed to reveal the temperature trend) (lines 211-212)

- Line 163: specify what do you mean with homogeneity and why you apply Pettitt test; response: more explanation was introduced (lines 222-223)

- Lines 163-167: why describe this test only with words and dedicate little less to a full page to Mann- Kendall test with all formulas?

response: Pettitt's test formulas are very long so the reference only is presented

- Line 181: Zc? response: corrected (line 245)

- Line 196: I think you are referring to Zc (attention when using the term p-value), and must declare it as the title of MK column in tables 2, 4, 6, 7. Furthermore, when you use *, **, *** in those tables I think that you have to clearly give meanings to these symbols in each caption. However, I understood what they mean, but they need an explicit explanation; response: the explicit explanation of all symbols was introduced (lines 311-338)

- Line 199-201:

why reporting global statistics before showing your results? They can have place in the Discussion section, and only if compared with local findings; response: moved to the appropriate place (lines 277-280)

- Lines 218-219: where? response: in many parts worldwide

- Lines 223-224: I think you should address the use of the word significant in the whole document, also in the light of considerations about test power response: the word significant was revised throughout the manuscript to assure that it used at the right place and used only when the significance level is found

Finally, we would like to thank you for your worthy comments that reinforced the paper totally

Response to Referee #2

Comments as shown in the comments on the manuscript PDF

1. be consistent with time (present or past) throughout the text response: the past time is used in the updated manuscript (line 10-22)

<u>*all line numbers correspond to the tracked changes version of the paper to facilitate</u> <u>revision</u>

2. are response: modified in the updated manuscript (line 19)

3. significant response: modified in the updated manuscript (line 22)

4. trends response: modified in the updated manuscript (line 22)

5. makes no sense

response: Blue Nile river only constitutes about 55% of the water reaches Egypt, and there are about 14 proposed dams and 5 agriculture projects. So, it is considered the most important and has the greatest potential among other basins in the River Nile Basin for development projects

6. such as response: modified in the updated manuscript (line 39)

7. the study of response: modified in the updated manuscript (line 42)

8. due to response: modified in the updated manuscript (line 46)

9. their response: modified in the updated manuscript (line 48)

10. result in response: modified in the updated manuscript (line 48)

11. can be noticed response: modified in the updated manuscript (line 54)

12. considered as response: modified in the updated manuscript (line 56)

13. and

response: modified in the updated manuscript (line 64)

14. New sentence: . In addition, higher peak flows have occurred in these basins response: modified in the updated manuscript (line 65)

15. Gleick (2000) response: modified in the updated manuscript (line 68)

16. almost

response: modified in the updated manuscript (line 72)

- 17. This has already been mentioned numerous times before. Generally, you have to better structure the introduction.
- 1) Why is climate change important (first generally and then for your basin)

2) Which variables need to be studied.

Why is temperature an important one.

3) What have other studies found regarding temperature trends in areas similar or close to yours.

4) What is the aim of the paper.

You have covered most of the points but you mix them up. response: modified in the updated manuscript (line 30-109)

18. MT

response: modified in the updated manuscript (line 79)

19. follow response: modified in the updated manuscript (line 80)

20. this sentence does not follow the previous response: this is the first sentence in a new paragraph (line 80)

21. (

response: modified in the updated manuscript (line 85)

22. (

response: modified in the updated manuscript (line 100)

23. need rephrasing

response: modified in the updated manuscript (line 100-101)

24. has been

response: modified in the updated manuscript (line 102)

25. the

response: modified in the updated manuscript (line 158)

26. ?

response: modified in the updated manuscript (line 160)

27. And

response: modified in the updated manuscript (line 161-162)

28. Could you also show the altitudes, and the lake and rivers position on the map? response: the map is modified totally in the updated manuscript (line 177)

29. which station? response: at Dangla station is inserted in the updated manuscript (line 151)

30. needs better phrasing.

Which are the technique categories that you used? I understand, from the next sentences, that The techniques are generally categorized as trend analysis and variability analysis techniques response: both techniques are used in the study, and the sentence was modified in the updated manuscript (line 185)

31. assuming

response: modified in the updated manuscript (line 196)

32. You mentioned the MK before. You should move all the information on the MK in one place. response: moved in the updated manuscript (line 227-230)

33. in

response: modified in the updated manuscript (line 293)

34. That response: modified in the updated manuscript (line 361)

35. thorough response: inserted in the updated manuscript (line 368)

1	Evaluation of Climate Change Impact on Extreme Temperature Variability in
2	the Blue Nile Basin, Ethiopia
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9	Abstract
10	Climate change raised important issues concerning the hydrological engineer. The impact of
11	climate change on important river basins should be investigated rigorously. Extreme temperature
12	variability has a direct impact on the hydrological cycle, especially the evaporation component.
13	In this paper, spatial and temporal patterns of changes in extreme temperatures are were
14	investigated using 10 meteorological stations data for the period 1950-2018 in the Blue Nile
15	basin. Monthly temperature data for the basin were used in the study. Long-term trends in the
16	Blue Nile Basin annual and monthly temperatures were investigated. The statistical significance
17	of the trend is-was calculated by applying the Mann-Kendall (MK) test. The analysis of data was
18	performed using the coefficient of variance, and anomaly index. The results showed that the
19	annual maximum and minimum temperatures is were increasing significantly with a magnitude
20	of 0.037° and 0.025° C per decade respectively in the period from 1950 to 2018. The result of the
21	Mann-Kendall analysis test revealed a marked increase in the mean maximum and minimum
22	temperatures trends over time significantly during the study period (the minimum temperature

rate is more evident than the maximum). The long-term anomalies of mean annual minimum
temperature revealed the inter-annual variability while the trend after 1977 was higher than the
long-term average that is proof of the warming trend's existence since the last two decades of the
20th century.

Keywords: Trend Analysis, Mann-Kendall Test, Climate Variability, Minimum Temperature,
Maximum Temperature, Blue Nile.

29 **1. Introduction**

30 Climate change is perceived to affect the whole world and there has been, and still, a growing 31 concern about its direction-trend and consequences. The Blue Nile River Basin (BNRB) is affected by climate change in temperature increase and rainfall decrease. The BNRB is the most 32 important and the potential river basins in the Nile Basin and having a great development 33 potential. Many development projects are under construction currently in the BNRB without 34 detailed climate change thorough analysis. These projects are essential for irrigation schemes, 35 36 hydropower generation, and national economic growth. As a trans-boundary river basin, it attracts the attention of some neighboring countries. For sound management and planning of 37 water resources, it is important to predict the climate change and variability of crucial 38 39 metrological factors such as temperature, and precipitation (Tamiru, 2011). Temperature and precipitation are affecting directly the hydrology of the basin (El-Mahdy et al., 2019). Since 40 41 temperature affects both evaporation and evapotranspiration, and precipitation is the source of 42 runoff, so the study of studying these factors is very important in water resources management-of the basin. The detailed analysis of the impact of these factors on basin hydrology will be 43 44 addressed by forthcoming researches. Here it is decided to explore the impact of climate change 45 on temperature and in a forthcoming paper the impact on rainfall will be studied.

The assessment of extreme temperature variability due to resulted from climate change will help 46 to better manage the development projects in the BNRB. Global and regional rises in air 47 48 temperature, together with their its collateral rises in water temperature, result in drive to negative alterations in water quality, even with the same precipitation (Collins et al., 2019). Lake 49 Tana, located upstream of BNRB, is one of the greatest natural lakes of Africa (El-Mahdy, 50 51 2014). Generally, the Lakes are sensitive to a wide range of climate change consequences. Even a slight change in climate can result in huge changes in lake salinity and levels (El-Mahdy et al., 52 53 2018, Mengistu et al., 2014, Tamiru, 2011).

54 The symptoms of global warming could can be noticed almost everywhere around the world. Heatwaves and droughts are striking a lot of places around the world, so that precipitations, 55 humidity, and temperatures are rarely normal. These phenomena among others are considered as 56 evidence of the presence of climate variability (Sohoulande Djebou and Singh, 2016). The 57 analysis of data from ground stations and satellites showed that the mean surface temperature of 58 the world has increased by about 0.6 degrees Celsius over the 20th century (Lindzen and 59 Giannitsis, 2002, Tapley et al., 2019). The warmest years ever have all occurred since 1980 in 60 this century (Janssens et al., 2020). The warmest recorded year was 1998 (Mann et al., 1999). 61 62 These alterations appear to be out of the natural variability range (Houghton, 1996, Mann et al., 1999, Sippel et al., 2020, Watson et al., 1996). A lot of researches studied the impact of 63 64 temperature rise on runoff, and it was found, with high confidence, that the runoff has increased in winter and decreased in summer and spring, __in-In_addition to higher peak flows have 65 66 occurred in these basins (Anache et al., 2018, Bergström et al., 2001, Chen et al., 2012, Yan et al., 2020). 67

The results of the simulation studies conducted by Gleick (2000) in the USA concluded that 68 69 small alterations in rainfall and temperature may lead to huge impacts on runoff. It was found that if temperature increases without a change in rainfall, the runoff will decrease. A 10% 70 decrease in rainfall together with a temperature rise of 2°C will reduce the runoff by 13-40%. On 71 the other hand, if rainfall increases by 10% and the temperature rise by 4°C, almost the rainfall 72 73 increase will balance evaporation losses and no alteration on runoff is predicted. Although these findings are not comprehensive, but they reveal the possible size and uncertainty encircling the 74 hydrologic and climatic implications of greenhouse warming (El-Mahdy, 2011). 75

76 Temperature is the most serious climate variable that affects hydrology and climate which, consequently, has an enormous direct and indirect impact on human life and the environment. 77 The temperature is a continuous variable in both space and time. The Monthly Mean 78 Temperature (MMT), as well as the maximum and minimum Monthly Temperature M(MT), 79 always follow matches a normal distribution (Jones and Hulme, 1996). (Jones and Hulme, 1996). 80 81 The predicted climate change may affect most of the Ethiopian districts. Drought remains Ethiopia's major hazard, while floods following in another place (Conway and Schipper, 2011). 82 Recently, it is clear that both drought and flood have increased in both frequency and size 83 84 (Margaret, 2003).

A recent study by Gebrehiwot and van der Veen (2013) indicated that climate change in Ethiopia
could lead to extreme temperatures and rainfall events, leading to more heavy and extended
droughts and floods. The climate of Ethiopia is composed of three distinct seasons named: Belg,
Kiremt, and Bega seasons (Haile, 2005). Belg is a short rain season which covers the period from
February to May. The rainfall in the Belg season results by the humid south-castern and eastern
winds that originate from the Indian Ocean. Kiremt is a long rainy season that starts in June and

normally ends in September. The Kiremt season's rainfall results substantially from the 91 convergence of low-pressure winds and the Intertropical Convergence Zone (ITCZ). Finally, the 92 dry season, called Bega, extends from October to January (Tabari et al., 2015). Changes in sea 93 surface temperature and El Niño Southern Oscillation (ENSO) periods in the Indian and Atlantic 94 Oceans have effects on the timing and amount of rainfall in Ethiopia (Fekadu, 2015, Shanko and 95 96 Camberlin, 1998). Asfaw et al. (2018) found that the drought events of Ethiopia are the result of both ENSO, Sea Surface Temperature (SST) variations in the Indian and Atlantic, and 97 anthropogenic activities. Kiros et al. (2016) found that the result of different studies of 98 temperature trend analyses in Northern Ethiopia is a mixture of non-significant negative and 99 positive trends. Mengistu et al. (2014) noted that the trend of temperature was significantly 100 101 increasing and statistically significant in Ethiopia. Although the concept of global warming and its impact on BNRB was has been studied before, but the analysis of long enough time series 102 103 over multiple weather stations with the appropriate statistical tools was not found. Investigation 104 of extreme temperature variability in the BNRB is not well presented in the literature. Making use of the statistical tools to explore climate change impact on Ethiopia is not performed in 105 thorough studies. The availability of the long time series data (1950-2018) invokes the research 106 107 on the climate change impact on the basin. The main objective of this research is to detect any climate change through the evaluation of extreme temperature variability using trend analysis of 108 109 along temperature time series in the BNRB, Ethiopia.

- 110 **2. Data and Methodology**
- 111 2.1.<u>1.1. Types of data</u>
- 112 Temperature data, which is used in the variability and time series trend analysis, has been
- 113 obtained from Ethiopia Meteorological Authority. Ten stations covering different parts of BNRB

114 have been chosen to study the variability of their data sets (see Table 1 and Figure 1). The elimate data consist of maximum & minimum MMT data probably with long years of records for 115 116 10 stations that have been collected in the BNRB. The length of record of the data made available for this analysis varies from 62 to 68 years, most analyses have focused on changes in 117 mean values due to the lack of the availability of high-quality daily data required for monitoring, 118 detecting, and attributing climate extremes changes (Jones, 1999). The study period (1950-2018) 119 was chosen according to the availability of the recorded data for all stations. The missing data 120 were calculated using the average of previous and subsequent months (for missed monthly data), 121 122 but the analysis excluded annual missing data. 123 124 125 126 127 Table (1): Selected weather stations of the Blue Nile basin and their general information. Altitude Mean annual Area (%) No Station Name Longitude (E) Latitude (N) rainfall (mm) (m) 4 Addis Ababa $\frac{2408}{2408}$ 1165 <u>38.7</u> 9.03 16.45 34.52 10.07 7.89 ₹ Assosa 1560 1126

3 Bahir Dar 37.42 11.06 1770 = = 4 **Debre Markos** 37.67 10.33 2515 1303 $\frac{11.84}{11.84}$ ₹ 37.04 2000 1102 Gondar 12.05 13.82 6 Gore 35.53 8.15 2002 8.55 2181 7 **Dangla** 36.8 $\frac{11.3}{11.3}$ 2030 4.61 1491 8 1480 Jimma 36.83 7.67 1676 $\frac{12.5}{12.5}$ 9.00 9 Sibu Sere 36.9 1750 8.55 1420 10 15.79 **Dessie** 39.67 11.08 2460 1045

128 (Source: Ethiopia Meteorological Authority)



upstream of BNRB, is the greatest natural lake in Ethiopia and the third greatest natural lake in

the Nile Basin, which is located in this basin. The main rivers that feed Lake Tana are Gilgil 141 Abai, Gomera, Rip, and Magek (El-Mahdy, 2017). These short-length rivers constitute 142 143 about 93% of the outflow of the lake (Kebede et al., 2006). The climate of Ethiopia is composed of three distinct seasons named: Belg, Kiremt, and Bega seasons (Haile, 2005). Belg is a short 144 rain season which covers the period from February to May. The rainfall in the Belg season 145 146 results by from the humid south-eastern and eastern winds that originate from the Indian Ocean. Kiremt is a long rainy season that starts in June and normally ends in September. The Kiremt 147 season's rainfall results substantially from the convergence of low-pressure winds and the 148 149 Intertropical Convergence Zone (ITCZ). Finally, the dry season, called the Bega, extends from October to January (Tabari et al., 2015). The highest temperature recorded in the BNRB is 150 151 located in the northwestern part at Dangla station. In some parts of Dabus, Dinder, Rihad, and 152 Beles, the maximum recorded temperature reaches 38°C and the minimum approaches 15°C. on 153 the other hand, the lower temperature is recorded in the eastern and central parts of the basin in 154 the Ethiopian highlands. In these areas, the maximum temperature is 20° C and the minimum is -155 1°C (Tamiru, 2011).

156 <u>2.2.Types of data</u>

157 <u>Temperature data, which is used in the variability and time series trend analysis, has been</u> 158 <u>obtained from the Ethiopia Meteorological Authority. Ten stations covering different parts of</u> 159 <u>BNRB have been chosen to study the variability of their data sets (see Table 1 and Figure 1). The</u> 160 <u>climate data consist of maximum & and minimum MMT data probably</u> with long years of records 161 <u>for 10 stations that have been collected in the BNRB. The length of the record of the data made</u> 162 <u>available for this analysis varies from 62 to 68 years of MMT, maximum, and minimum MT₃.</u> 163 <u>Most analyses have focused on changes in mean values due to the lack of the availability of</u>

164	<u>high-q</u>	<u>uality daily data</u>	required for m	onitoring, dete	cting, and	attributing o	climate extremes						
165	change	es (Jones, 1999). 7	The study period	<u>(1950-2018) w</u>	as chosen ac	ccording to t	<u>he availability of</u>						
166	the rec	corded data for all	stations. Althoug	gh the missing of	lata for the s	tudied perio	d is almost none.						
167	But for the missing data of a specific month werewas calculated using the average of previous												
168	year data of the same month and subsequent year months data (for missed monthly data), but the												
169	analysis excluded annual missing data.												
170													
171													
172													
173													
174	Table	(1): Selected weat	ther stations of th	e Blue Nile bas	in and their	general info	<u>rmation.</u>						
	<u>No</u>	Station Name	Longitude (E)	Latitude (N)	<u>Altitude</u> (m)	<u>Area (%)</u>	<u>Mean annual</u> rainfall (mm)						
	1	Addis Ababa	38.7	9.03	2408	16.45	1165						
	2	Assosa	34.52	10.07	1560	7.89	1126						
	<u>3</u>	Bahir Dar	37.42	11.06	1770	-	-						
	4	Debre Markos	37.67	<u>10.33</u>	2515	<u>11.84</u>	<u>1303</u>						

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<u>6</u>

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10Dessie39.675(Source: Ethiopia Meteorological Authority)

<u>37.04</u>

<u>35.53</u>

36.8

36.83

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Gondar

Gore

<u>Dangla</u>

Jimma

Sibu Sere

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<u>11.3</u>

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<u>2030</u>

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183 **2.3. Methods**

184 The analysis of temperature was performed using many techniques. These techniques are 185 generally categorized as variability and trend analysis and variability as standard deviation (SD). The variability analysis encompasses the Coefficient of Variation (CV), moving average, and the 186 187 percentage departure from the mean (Gebrechorkos et al., 2020). On the other hand, trend 188 analysis is applied to consistent data only utilizing parametric and non-parametric tests 189 (Onyutha, 2017, Partal and Kahya, 2006). The homogeneity and normality of the variance over 190 the data series are usually affected by missing data and outliers in parametric tests. The nonparametric test is used to overcome the problem of non-normally distributed, missing data, and 191 outliers. This problem is repeatedly found in the hydrological and meteorological time series. For 192 these reasons, Mann- Kendall (MK) test is commonly utilized to discover the trends of 193 meteorological and hydrological variables (Degefu and Bewket, 2014, Kiros et al., 2016, Seleshi 194 195 and Camberlin, 2006, Tabari and Tavakoli, 2016). MK is a nonparametric test, that assesses the 196 trend of any time series without assuming assigning if the trend is non-linear or linear. MK test is commonly used to discover the monotonic (increasing or decreasing) trends in the time series of 197 198 hydrological, meteorological, and environmental data (Yue et al., 2002). Totaro et al. (2020) 199 conducted trend detection using a numerical investigation. The study focused on exploring the 200 power of nonparametric and parametric tests in annual maximum time series data. The results 201 showed paramount practical implications. It was proved that the dependence of test power on the parameters of parent distribution might affect the results significantly of both nonparametric and 202 203 parametric tests. This result is comprising the widely applied MK test. The results of Monte 204 Carlo simulations showed that the power of the MK test has a direct relationship with the data

205 variance in addition to the trend magnitude (Wang et al., 2020). Németh et al. (2020) analyzed the characteristics of the test of the likelihood ratio for extremes by simulations and introduced a 206 simulation-based method to overcome the issue of scarce data. A novel return level calculation 207 procedure is -proposed. The probability or power of discovering trends relies on the development 208 209 of efficient multivariate statistical and deterministic procedures for forecasting future trends in 210 processes of the earth system (Vogel et al., 2013). In the current paper, the variability of temperature was calculated employing the SD and CV tests. Moreover, MK and Sen's slope 211 212 estimator test were employed to reveal the temperature trend Moreover, MK was employed to 213 reveal the temperature trend using Sen's slope estimator. SPSS v22 software was used to perform the data analysis. CV is computed to evaluate the rainfall variability. the higher the CV value, the 214 larger variability of rainfall. CV is calculated using the following equation: 215

$$216 \qquad CV = \frac{\sigma}{\mu} \times 100 \tag{1}$$

217 Where CV is the coefficient of variation; σ is the standard deviation and μ is the mean (Isioma et al., 2018). MK test was employed to explore the presence of monotonic trends in the time series, 218 and the statistical significance of the trend. Since outliers cannot be averted in any time-series, 219 220 the MK test is advantageous because it is based on the (+ or -) signs instead of the random 221 variable values. Therefore, the MK test is less influenced by the outliers (Asfaw et al., 2018). 222 The test of homogeneity was conducted employing Pettitt's test (Pettitt, 1979) to ensure that the 223 data are homogenous and no misleading data are present. In this test, every value in the time 224 series is compared with all posterior values in the data series. If the value of the subsequent 225 period is greater than the value of the preceding period, the statistic (S) is raised by one and vice versa. The summation of all increments and decrements reveals the total value of the statistic (S). 226 227 MK test was employed to explore the presence of monotonic trends in the time series, and the

<u>statistical significance of the trend. Since outliers cannot be averted in any time-series, the MK</u>
 <u>test is advantageous because it is based on the (+ or -) signs instead of the random variable</u>
 <u>values. Therefore, the MK test is less influenced by the outliers (Asfaw et al., 2018).</u> The MK
 test statistic 'S' is calculated based on (Mann, 1945, McLeod, 2005, Yue et al., 2002), using the
 formula:

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$$S = \sum_{i=1}^{n=1} \sum_{j=i+1}^{n} sgn(x_j - x_i)$$
(2)

The trend test is applied on a time series x_i which is ranked as i = 1, 2..., n-1 and x_j , that is ranked as j = i - 1, 2..., n. Each of the data point x_i is taken as a reference point to be compared with the other data point's x_i so that:

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$$Sgn(x_j - x_i) = \begin{cases} +1 \ if \ (x_j - x_i) > 0 \\ 0 \ if \ (x_j - x_i) = 0 \\ -1 \ if \ (x_j - x_i) < 0 \end{cases}$$

where x_i and x_j are the data values in the years i and j respectively (j > i).

It was reported that, if the number of observations is greater than 10 ($n \ge 10$), the statistic (S) is usually normally distributed around the mean (McLeod, 2005). So that, the variance could be formulated as:

242
$$Var(S) = \frac{n(n-1)(2n+5) - \sum_{i=1}^{m} t_1(t_1-1)(2t_1+5)}{18}$$
 (4)

where n is the number of observation and t_i are the ties of the sample time series. The test statistics Z_c is as follows:

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$$Z_c Z = \begin{cases} \frac{S-1}{\sigma} if & S > 0\\ 0 & if & S = 0(5)\\ \frac{S+1}{\sigma} if & S < 0 \end{cases}$$

If Z_c follows a normal distribution, its positive or negative value refers to an upward or downward trend for the studied period respectively. The trend magnitude could be calculated by slope estimator methods (Sen, 1968). Here, the slope (T_i) of all data pairs is calculated according to (Sen, 1968). In general, T_i for any time series x could be predicted from:

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$$T_i = \frac{x_j - x_k}{j - k}$$
 (6)

where x_j and x_k are considered as data values at time j and k (j > i) correspondingly. The median of these N values of T_i is represented as Sen's estimator of the slope which is calculated as $Q_{med}=T_{(N+1)/2}$ if N is an odd number, and it is computed as $Q_{med}=[T_{N/2} + T(_{(N+2)/2})/2]$ if N is an even number. A positive value of T_i indicates an upward or increasing trend and a negative value of T_i gives a downward or decreasing trend in the time series.

256 **3. Results and Discussion**

The MK test and Sen's slope estimator were applied to the time-series 1950–2018 for the ten meteorological stations in BNRB. The annual and seasonal mean of the climatic parameters, particularly minimum and maximum MMT were analyzed. Table 2 shows the MK test statistics and p values at 5%, 1%, and 0.1% levels of significances. In the MK test, parameters such as Kendall's tau, S, and Z statistics were computed to determine the positive or negative trend of climate parameters in the long time series studied.

Surface air temperature rises globally, although global warming is not uniform over the world, neither temporally nor spatially (Turner et al., 2020). The surface temperature increase is the direct demonstration of global climate change (Brunet et al., 2005). Tables 2 and 4 showed the annual and monthly <u>maximum and minimum MT and MMT maximum and minimum</u> and their trend during the study period. The mean minimum temperature in the BNRB area was 6.7°C and the mean maximum was 25.5°C, with a mean annual temperature of 16.1°C. The regression line

slope is about 0.036° and 0.024°C per decade for minimum and maximum annual temperatures 269 respectively for the studied period of 1950-2018 as shown in Tables 3 and 5 in addition to 270 Figures 2 and 3. These results are close to the previously found global warming rate which is 271 0.06° C per decade for the last century (Pachauri and Meyer, 2014). Figures 4 and 5 showed the 272 long-range anomalies of mean annual minimum and maximum temperatures. It is clear that the 273 274 trend after 1985 was greater than the long-term average. This is considered as strong evidence for the global warming trend since the 20th century last two decades. This result enhances the 275 276 previously found results that the climate becames warmer since the end of the last century 277 (Bathiany et al., 2018, Guo et al., 2020, Nijsse et al., 2019). Surface air temperature rises globally, although global warming is not uniform over the world, neither temporally nor spatially 278 279 (Turner et al., 2020). The surface temperature increase is the direct demonstration of global climate change (Brunet et al., 2005). However, other researchers proved that the last four 280 decades of the 20th century showed a significantly increasing trend in global warming 281 282 (Bhutiyani et al., 2008). Interannual variability of rainfall and temperature has intensified from the late 1960s. Also, droughts occur more frequently and over a greater spatial distribution. A 283 stronger statistical relation between climate change and El Niño-Southern Oscillation (ENSO) 284 285 phenomenon is observed (Fauchereau et al., 2003).

As shown in Tables 2 and 4, the result of the MK test detected that the mean maximum and minimum temperatures increased significantly overtime at a 99% confidence level. The trend for monthly minimum and maximum temperature is increasing significantly for all months. The gross increase in the observed annual temperature is attributed to the increase in the minimum temperature. Many researchers have reported that the minimum temperature rises more than the maximum (Bayramzadeh et al., 2018, Crimp et al., 2018, Gross et al., 2018, Scott et al., 2017)

292 The minimum temperature incremental increase is more evident than that of the maximum. The 293 results are of in good agreement with the results by (Asfaw et al., 2018). The results of the MK 294 analysis revealed an increase in both mean and minimum temperatures throughout the studied period significantly. On the other hand, the maximum temperature has a non-significant 295 increasing trend. The studies that agree with these results such as (Mengistu et al., 2014, Roy and 296 297 Das, 2013) found that the increasing trends in the maximum temperature series were less than the 298 minimum temperature series. The cause that the minimum temperature increases more than the maximum may be explained as the sensitivity of minimum temperature to the greenhouse is 299 300 higher than the maximum (Salawitch, 1998). In this regard, the maximum temperature over India was higher than the mean for the period (1901–2007) with a low trend. However, the minimum 301 302 temperature displayed an upward trend (Kothawale et al., 2010). This result supported the findings of the current paper over BNRB. Table 6 showed that the minimum temperature has an 303 increasing trend in Bega, Belg, and Kiremt seasons in the stations of Bahir Dar, Debre Markos, 304 305 Gondar, Gore, and Jimma with a level of significance 5%, 1%, and 0.1%. Also, the maximum temperature in the three seasons has a significant increasing trend in Addis Ababa, Bahir Dar, 306 Gondar, and Gore stations and the trend is statistically significant at 0.1%, 1%, and 5% level of 307 308 significance as shown in Table 7.

Month	Min	Mor	Moon	۲D	CV(0/)	Skewness/	MK	Sens
Month	MIII.	Max.	Mean	2D	CV(%)	Kurtosis	Test	slope
January	6.7	11.8	10.0	1.2	11.8	-0.9/0.7	5.26***	0.03
February	8.6	13.5	11.3	1.1	10.0	-0.5/-0.1	5.62***	0.03
March	10.0	14.6	12.8	1.1	8.8	-0.8/-0.4	5.70***	0.03
April	11.2	15.5	13.6	1.0	7.4	-0.5/-0.7	5.54***	0.03
May	11.6	15.4	13.8	1.0	7.2	-0.5/-0.8	5.63***	0.03
June	10.7	15.0	13.6	1.1	7.8	-0.9/-0.3	5.75***	0.03
July	10.7	14.9	13.5	0.9	6.9	-0.8/-0.1	6.76***	0.03

Table (2): Basic statistics and MK trend analysis of minimum temperature in the Blue Nile basin
 (1950–2018).

August	10.5	14.7	13.3	0.9	6.8	-0.8/0.0	7.45***	0.039
September	10.0	14.3	12.9	0.9	7.0	-0.8/0.1	6.83***	0.035
October	8.8	13.1	11.5	1.0	9.1	-0.7/-0.5	6.73***	0.040
November	8.0	12.3	10.2	1.0	9.5	-0.5/-0.5	5.24***	0.031
December	7.1	11.9	9.5	1.1	11.7	-0.1/-0.8	5.47***	0.036
Annual	10.3	13.4	12.2	0.9	7.5	-0.7/-0.9	6.53***	0.036
Bega (ONDJ)	8.2	11.8	10.3	0.9	8.9	-0.6/-0.8	6.85***	0.035
Belg (FMAM)	10.6	14.5	12.9	1.0	7.9	-0.6/-0.7	5.70***	0.036
Kiremt (JJAS)	10.4	14.6	13.3	0.9	7.0	-0.9/0.0	6.96***	0.037

311 <u>* Significant at 5% significance level</u>

312 <u>** Significant at 1% significance level</u>

313 <u>*** Significant at 0.1% significance level</u>

Table (3): Linear regression result (Annual minimum temperature, Bega, Belg and Kiremt season) (1950–2018)

Season	Change in minimum (°C/year)	P-value	\mathbb{R}^2	CV (%)
Annual	0.037	0.001	0.65	7.5
Bega (ONDJ)	0.037	0.001	0.64	8.9
Belg (FMAM)	0.038	0.001	0.55	7.9
Kiremt (JJAS)	0.037	0.001	0.63	7.0

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- Table (4): Basic statistics and MK trend analysis of maximum temperature in the Blue Nile basin(1950–2018).

Month	Min	Mov	Maan	SD	CV(0(1))	Skewness/	MK	Sens
IVIOIIIII	IVIIII.	IVIAA.	wican	3D	CV(%)	Kurtosis	Test	slope
January	24.5	27.9	26.5	0.8	2.9	-0.6/-0.3	4.88***	0.023
February	24.2	30.0	27.6	1.1	3.8	-0.5/1.1	4.93***	0.027
March	25.5	30.9	28.4	1.0	3.6	-0.6/0.9	5.39***	0.027
April	25.5	30.3	28.0	1.1	4.0	-0.4/-0.6	5.51***	0.036
May	25.2	29.0	27.1	0.9	3.2	-0.3/-0.6	4.59***	0.024
June	23.3	26.6	25.1	0.7	2.9	-0.5/-0.3	5.13***	0.022
July	20.8	25.0	23.2	0.8	3.5	-0.5/0.4	5.94***	0.024
August	21.3	24.2	23.1	0.7	3.0	-0.6/-0.4	6.14***	0.024

September	22.6	25.4	24.2	0.6	2.5	-0.4/-0.1	6.06***	0.019
October	23.3	26.2	25.1	0.6	2.6	-0.5/0.0	4.78***	0.018
November	23.5	26.8	25.6	0.8	3.1	-0.7/-0.3	5.26***	0.022
December	23.2	26.9	25.9	0.8	3.1	-1.1/0.9	5.82***	0.025
Annual	24.5	26.8	25.8	0.6	2.5	-0.5/-1.0	7.09***	0.024
Bega (ONDJ)	24.1	26.8	25.8	0.7	2.6	-0.7/-0.5	6.05***	0.023
Belg (FMAM)	25.6	29.2	27.8	0.9	3.1	-0.5/-0.2	6.16***	0.028
Kiremt (JJAS)	22.4	25.0	23.9	0.6	2.6	-0.8/-0.1	6.59***	0.021

* Significant at 5% significance level ** Significant at 1% significance level *** Significant at 0.1% significance level

Table (5): Linear regression result (Annual maximum temperature, Bega, Belg and Kiremt season) (1950–2018)

Season	Change in maximum (°C/year)	P-value	\mathbb{R}^2	CV (%)
Annual	0.025	0.001	0.644	2.5
Bega (ONDJ)	0.023	0.001	0.493	2.6
Belg (FMAM)	0.030	0.001	0.482	3.1
Kiremt (JJAS)	0.023	0.001	0.534	2.6

	Season		Addis Ababa	Assosa	Bahir Dar	Debre Markos	Gondar	Gore	Dangla	Jimma	Sibu Sere	Dessie
		Mean	9.5	15.1	11.9	9.6	13.0	13.5	18.3	11.6	11.0	11.9
		SD	0.9	0.7	1.4	0.9	0.7	0.5	0.8	1.0	0.5	1.3
	Annual	CV	0.1	0.0	0.1	0.1	0.1	0.0	0.0	0.1	0.0	0.1
		MK	1.22	0.3	5.1***	7.0***	4.5***	4.4***	-2.6*	4.5***	2.7**	1.2
		Sen's slope	0.005	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
		Mean	7.6	14.2	9.8	8.2	11.6	13.5	15.5	8.3	8.4	9.4
ature (°c)		SD	1.2	0.9	1.6	1.0	0.9	0.5	1.3	1.6	1.0	1.1
	Bega	CV	0.2	0.1	0.2	0.1	0.1	0.0	0.1	0.2	0.1	0.1
		MK	0.33	-0.3	4.9***	6.7***	4.8***	4.1***	-1.8	4.6***	0.9	1.0
pera		Sen's slope	0.002	0.0	0.1	0.0	0.0	0.0	0.0	0.0	0.0	0.0
luia												
E E		Mean	10.5	16.1	12.4	10.5	14.3	14.2	18.8	11.7	11.5	12.5
mu		SD	0.9	1.2	1.7	1.1	0.9	0.7	1.0	1.5	0.8	1.5
ini	Belg	CV	0.1	0.1	0.1	0.1	0.1	0.0	0.1	0.1	0.1	0.1
Z		MK	0.99	0.6	5.0***	5.9***	3.7***	2.7**	-2.7**	3.2**	3.1**	1.2
		Sen's slope	0.006	0.0	0.1	0.0	0.0	0.0	0.0	0.0	0.0	0.0
		Mean	10.6	15.1	13.7	10.0	13.2	12.8	20.5	14.6	13.1	13.9
		SD	0.7	0.8	1.2	0.9	0.7	0.4	0.6	1.2	0.4	1.8
	Kiremt	CV	0.1	0.1	0.1	0.1	0.1	0.0	0.0	0.1	0.0	0.1
		MK	2.17*	1.4	4.7***	6.9***	3.4***	5.6***	-2.0*	2.7**	2.2*	1.5
		Sen's slope	0.01	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0

Table (6): Mann-Kendal minimum temperature trend analysis (Based on gauge stations)

* Significant at 5% significance level

332 <u>** Significant at 1% significance level</u>

333 <u>*** Significant at 0.1% significance level</u>

Table (7): Mann-Kendal maximum temperature trend analysis (Based on gauge stations)

	Season		Addis Ababa	Assosa	Bahir	Debre Markos	Gondar	Gore	Dangla	Jimma	Sibu Soro	Dessie
		Mean	Ababa 22.8	28.2	26.8	22.4	26.6	23.5	29.8	27.3	26.9	26.2
		SD	0.8	0.6	0.6	0.4	0.5	0.5	1.8	0.5	0.5	0.7
	Annual	CV	0.0	0.0	0.0	0.0	0.0	0.0	0.1	0.0	0.0	0.0
		MK	5.25***	1.6	4.3***	3.4***	4.0***	4.8***	1.5	1.5	-0.2	2.5*
		Sen's slope	0.026	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
tture (°c)		-										
		Mean	22.7	28.6	26.4	22.7	26.9	23.9	29.7	27.2	27.5	24.4
		SD	1.0	0.7	0.7	0.7	0.6	0.6	1.8	0.4	0.6	0.8
	Bega	CV	0.0	0.0	0.0	0.0	0.0	0.0	0.1	0.0	0.0	0.0
oer:		MK	5.08***	2.1*	4.4***	3.1**	3.5***	3.9***	1.1	-0.4	0.6	4.2***
fme		Sen's slope	0.027	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
ц Ц												
mm		Mean	24.4	31.2	29.1	24.7	29.0	25.4	33.5	28.4	28.4	26.7
axi		SD	1.2	0.8	0.9	0.7	0.9	0.7	2.1	0.8	0.7	1.0
Σ	Belg	CV	0.0	0.0	0.0	0.0	0.0	0.0	0.1	0.0	0.0	0.0
		MK	3.44***	1.8	2.5*	2.3*	2.4*	2.7**	2.4*	1.5	-0.6	0.2
		Sen's slope	0.026	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
	-											
		Mean	21.3	25.0	24.9	19.7	23.9	21.3	26.2	26.5	24.8	27.6
	Kiremt	SD	0.8	0.7	0.6	0.4	0.6	0.5	2.0	0.6	0.6	0.7
		CV	0.0	0.0	0.0	0.0	0.0	0.0	0.1	0.0	0.0	0.0

МК	4.41***	0.2	3.9***	1.1	2.8**	4.5***	1.3	1.8	-0.6	1.8
Sen's slope	0.021	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0

336 <u>* Significant at 5% significance level</u>

337 <u>** Significant at 1% significance level</u>

338 <u>*** Significant at 0.1% significance level</u>





Figure (2): Linear regression result of annual minimum temperature (1950–2018)





Figure (3): Linear regression result of annual maximum temperature (1950–2018)







Temporal and spatial patterns of changes in extreme temperatures are investigated in the Blue Nile River basin. The study used data from 10 meteorological stations in the basin for the period 1950-2018. The statistical data analysis used in this research work includes standard deviation, coefficient of variation, Skewness/ Kurtosis, Mann-Kendall Test, and Sens slope. The results showed that there is an increase in both maximum and minimum temperatures in the basin over the study period of 68 years. The rate of temperature

356 increase was found to be 0.037° and 0.025°C per decade for minimum and maximum temperatures, respectively. The trend of the mean minimum annual temperature for the 357 last quarter of the 20th century was greater than the long-term average. The trend analysis 358 test showed a significant increase of the mean minimum and maximum temperatures for 359 all months. The rate of increase for the minimum temperature is more pronounced than 360 361 the maximum. Consequently, this indicates the that minimum temperature is more likely sensitive to climate change. Some stations showed a decrease in the inter-annual 362 363 temperature range. rational climate change may aggravate the situations of climate 364 extremes. Therefore, appropriate adaptation and mitigation strategies should be planned to lessen the impacts of such climatic risks in the Blue Nile River Basin. The detailed 365 analysis of the impact of these factors on basin hydrology will be addressed by 366 forthcoming researches. A further enhancement of the current work could be attained 367 using a long record of daily data. More A more thorough investigation of the techniques 368 369 used for trend analysis could lead to more accurate trend values. A Holistic climate change assessment over the whole Nile River Basin should get more attention and a 370 371 thorough investigation. Researches on climate change detection should be done in various 372 approaches in collaboration with worldwide concerned bodies.

5. Data Availability Statement

Some or all data, models, or code generated or used during the study are available fromthe corresponding author by request

6. Author contribution

All the process of the paper including but not limited to: Conceptualization, Datacuration, Formal analysis, Investigation, Methodology, Project administration, Resources,

- 379 Software, Supervision, Validation, Visualization, Writing original draft preparation,
- and Writing review & editing were done by the Mostafa A. Mohamed and Mohamed
- 381 El-Sayed El-Mahdy in equal share

382 7. Competing interests

383 The authors declare that they have no conflict of interest.

384 **8. References**

- ANACHE, J. A., FLANAGAN, D. C., SRIVASTAVA, A. & WENDLAND, E. C. 2018. Land use and
 climate change impacts on runoff and soil erosion at the hillslope scale in the Brazilian
 Cerrado. Science of the Total Environment, 622, 140-151.
- ASFAW, A., SIMANE, B., HASSEN, A. & BANTIDER, A. 2018. Variability and time series trend
 analysis of rainfall and temperature in northcentral Ethiopia: A case study in Woleka
 sub-basin. *Weather and climate extremes*, 19, 29-41.
- 391 BATHIANY, S., DAKOS, V., SCHEFFER, M. & LENTON, T. M. 2018. Climate models predict 392 increasing temperature variability in poor countries. *Science advances*, 4, eaar5809.
- BAYRAMZADEH, V., ZHU, H., LU, X., ATTAROD, P., ZHANG, H., LI, X., ASAD, F. & LIANG, E. 2018.
 Temperature variability in northern Iran during the past 700 years. *Sci. Bull*, 63, 462-464.
- BERGSTRÖM, S., CARLSSON, B., GARDELIN, M., LINDSTRÖM, G., PETTERSSON, A. &
 RUMMUKAINEN, M. 2001. Climate change impacts on runoff in Sweden assessments by
 global climate models, dynamical downscaling and hydrological modelling. *Climate research*, 16, 101-112.
- BHUTIYANI, M., KALE, V. S. & PAWAR, N. 2008. Changing streamflow patterns in the rivers of
 northwestern Himalaya: implications of global warming in the 20th century. *Current Science*, 618-626.
- BRUNET, M., JONES, P., SIGRÓ, J., SALADIÉ, O., AGUILAR, E., MOBERG, A., DELLA-MARTA, P.,
 LISTER, D., WALTHER, A. & LÓPEZ, D. 2005. Spatial and temporal temperature variability
 and change over Spain during 1850-2003. *Journal of Geophysical Research Atmospheres*.
- CHEN, H., XU, C.-Y. & GUO, S. 2012. Comparison and evaluation of multiple GCMs, statistical downscaling and hydrological models in the study of climate change impacts on runoff.
 Journal of hydrology, 434, 36-45.
- 409 COLLINS, S., YUAN, S., TAN, P., OLIVER, S., LAPIERRE, J., CHERUVELIL, K., FERGUS, C., SKAFF, N.,
 410 STACHELEK, J. & WAGNER, T. 2019. Winter precipitation and summer temperature
 411 predict lake water quality at macroscales. *Water Resources Research*, 55, 2708-2721.
- 412 CONWAY, D. & SCHIPPER, E. L. F. 2011. Adaptation to climate change in Africa: Challenges and 413 opportunities identified from Ethiopia. *Global Environmental Change*, 21, 227-237.
- CRIMP, S., NICHOLLS, N., KOKIC, P., RISBEY, J. S., GOBBETT, D. & HOWDEN, M. 2018. Synoptic to
 large-scale drivers of minimum temperature variability in Australia–long-term changes.
 International Journal of Climatology, 38, e237-e254.
- DEGEFU, M. A. & BEWKET, W. 2014. Variability and trends in rainfall amount and extreme event
 indices in the Omo-Ghibe River Basin, Ethiopia. *Regional environmental change*, 14, 799810.

- 420 EL-MAHDY, M. E.-S., ABBAS, M. S. & SOBHY, H. M. 2019. Development of mass-transfer 421 evaporation model for Lake Nasser, Egypt. *Journal of Water and Climate Change*.
- 422 EL-MAHDY, M. E. 2011. *Recent changes of Lake Nasser hydrology*. MSc. thesis, P241, Cairo 423 University.
- 424 EL-MAHDY, M. E. Tana Lake Free Water Evaporation. Sustainable Development of Natural 425 Resources in the Nile Basin Countries-Cairo University, Giza, Egypt, 2014. 353-389.
- 426 EL-MAHDY, M. E. 2017. *Modelling Eastern Nile Basin Hydrology*. Cairo University.
- 427 EL-MAHDY, M. E., ABBAS, M. S. & SOBHY, H. M. 2018. Investigating the Water Quality of the 428 Water Resources Bank of Egypt: Lake Nasser. *Conventional Water Resources and* 429 *Agriculture in Egypt.* Springer.
- FAUCHEREAU, N., TRZASKA, S., ROUAULT, M. & RICHARD, Y. 2003. Rainfall variability and
 changes in southern Africa during the 20th century in the global warming context. *Natural hazards*, 29, 139-154.
- 433 FEKADU, K. 2015. Ethiopian seasonal rainfall variability and prediction using canonical 434 correlation analysis (CCA). *Earth Sciences*, 4, 112.
- GEBRECHORKOS, S. H., HÜLSMANN, S. & BERNHOFER, C. 2020. Analysis of climate variability and
 droughts in East Africa using high-resolution climate data products. *Global and Planetary Change*, 186, 103130.
- GEBREHIWOT, T. & VAN DER VEEN, A. 2013. Assessing the evidence of climate variability in the
 northern part of Ethiopia. *Journal of development and agricultural economics*, 5, 104 119.
- GLEICK, P. H. 2000. Water: the potential consequences of climate variability and change for the
 water resources of the United States, Pacific Institute for Studies in Development,
 Environment, and Security.
- GROSS, M. H., DONAT, M. G., ALEXANDER, L. V. & SISSON, S. A. 2018. The sensitivity of daily
 temperature variability and extremes to dataset choice. *Journal of Climate*, 31, 13371359.
- 447 GUO, Y., LI, Y., WANG, F., WEI, Y. & RONG, Z. 2020. Processes Controlling Sea Surface 448 Temperature Variability of Ningaloo Niño. *Journal of Climate*, 33, 4369-4389.
- HAILE, M. 2005. Weather patterns, food security and humanitarian response in sub-Saharan
 Africa. *Philosophical Transactions of the Royal Society of London B: Biological Sciences*,
 360, 2169-2182.
- HOUGHTON, E. 1996. Climate change 1995: The science of climate change: contribution of
 working group I to the second assessment report of the Intergovernmental Panel on
 Climate Change, Cambridge University Press.
- ISIOMA, I. N., RUDOLPH, I. I. & OMENA, A. L. 2018. Non-parametric Mann-Kendall Test Statistics
 for Rainfall Trend Analysis in Some Selected States within the Coastal Region of Nigeria.
 Journal of Civil, Construction and Environmental Engineering, 3, 17.
- JANSSENS, C., HAVLÍK, P., KRISZTIN, T., BAKER, J., FRANK, S., HASEGAWA, T., LECLÈRE, D., OHREL,
 S., RAGNAUTH, S. & SCHMID, E. 2020. Global hunger and climate change adaptation
 through international trade. *Nature Climate Change*, 1-7.
- JONES, P. 1999. Sur face Air Tem per a ture and Its Vari a tions over the Last 150 Years. *Rev. Geophys*, 37.
- JONES, P. D. & HULME, M. 1996. Calculating regional climatic time series for temperature and
 precipitation: methods and illustrations. *International Journal of Climatology: A Journal*of the Royal Meteorological Society, 16, 361-377.

- KEBEDE, S., TRAVI, Y., ALEMAYEHU, T. & MARC, V. 2006. Water balance of Lake Tana and its
 sensitivity to fluctuations in rainfall, Blue Nile basin, Ethiopia. *Journal of hydrology*, 316,
 233-247.
- 469 KIROS, G., SHETTY, A. & NANDAGIRI, L. 2016. Analysis of variability and trends in rainfall over 470 northern Ethiopia. *Arabian Journal of Geosciences*, 9, 451.
- KOTHAWALE, D., MUNOT, A. & KUMAR, K. K. 2010. Surface air temperature variability over India
 during 1901–2007, and its association with ENSO. *Climate Research*, 42, 89-104.
- LINDZEN, R. S. & GIANNITSIS, C. 2002. Reconciling observations of global temperature change.
 Geophysical research letters, 29, 24-1-24-3.
- 475 MANN, H. B. 1945. Nonparametric tests against trend. *Econometrica: Journal of the Econometric* 476 *Society*, 245-259.
- 477 MANN, M. E., BRADLEY, R. S. & HUGHES, M. K. 1999. Northern hemisphere temperatures during
 478 the past millennium: Inferences, uncertainties, and limitations. *Geophysical research*479 *letters*, 26, 759-762.

480 MARGARET, F. 2003. Planning for the Next Drought, Ethiopia Case Study. USAID, Washington.

- 481 MCLEOD, A. I. 2005. Kendall rank correlation and Mann-Kendall trend test. *R Package Kendall*.
- MENGISTU, D., BEWKET, W. & LAL, R. 2014. Recent spatiotemporal temperature and rainfall
 variability and trends over the Upper Blue Nile River Basin, Ethiopia. International
 Journal of Climatology, 34, 2278-2292.
- 485 NÉMETH, L., HÜBNEROVÁ, Z. & ZEMPLÉNI, A. 2020. Comparison of trend detection methods in 486 GEV models. *Communications in Statistics-Simulation and Computation*, 1-16.
- NIJSSE, F. J., COX, P. M., HUNTINGFORD, C. & WILLIAMSON, M. S. 2019. Decadal global
 temperature variability increases strongly with climate sensitivity. *Nature Climate Change*, 9, 598-601.
- 490 ONYUTHA, C. 2017. On rigorous drought assessment using daily time scale: Non-stationary
 491 frequency analyses, revisited concepts, and a new method to yield non-parametric
 492 indices. *Hydrology*, 4, 48.
- PACHAURI, R. K. & MEYER, L. 2014. Climate change 2014 Synthesis Report-Summary for
 Policymakers. Intergovernmetnal Panel on Climate Change (IPCC).
- PARTAL, T. & KAHYA, E. 2006. Trend analysis in Turkish precipitation data. *Hydrological Processes: An International Journal*, 20, 2011-2026.
- 497 PETTITT, A. 1979. A non-parametric approach to the change-point problem. *Journal of the Royal* 498 Statistical Society: Series C (Applied Statistics), 28, 126-135.
- ROY, T. D. & DAS, K. K. 2013. Temperature Trends at Four Stations of Assam during the period
 1981-2010. Int. J. Sci. Res. Publ, 3, 1-3.
- 501 SALAWITCH, R. J. 1998. A greenhouse warming connection. *Nature*, 392, 551-552.
- 502 SCOTT, A. A., ZAITCHIK, B., WAUGH, D. W. & O'MEARA, K. 2017. Intraurban temperature 503 variability in Baltimore. *Journal of Applied Meteorology and Climatology*, 56, 159-171.
- 504 SELESHI, Y. & CAMBERLIN, P. 2006. Recent changes in dry spell and extreme rainfall events in 505 Ethiopia. *Theoretical and Applied Climatology*, 83, 181-191.
- 506 SEN, P. K. 1968. Estimates of the regression coefficient based on Kendall's tau. *Journal of the* 507 *American statistical association,* 63, 1379-1389.
- SHANKO, D. & CAMBERLIN, P. 1998. The effects of the Southwest Indian Ocean tropical cyclones
 on Ethiopian drought. *International Journal of Climatology*, 18, 1373-1388.
- 510 SINTAYEHU, L. 2015. Application of the HEC-HMS model for runoff simulation of upper blue Nile 511 River Basin. *Hydrol Current Res*, 6, 1-8.

- SIPPEL, S., MEINSHAUSEN, N., FISCHER, E. M., SZÉKELY, E. & KNUTTI, R. 2020. Climate change
 now detectable from any single day of weather at global scale. *Nature Climate Change*,
 10, 35-41.
- 515 SOHOULANDE DJEBOU, D. C. & SINGH, V. P. 2016. Impact of climate change on precipitation 516 patterns: A comparative approach. *International Journal of Climatology*, 36, 3588-3606.
- TABARI, H., TAYE, M. T. & WILLEMS, P. 2015. Statistical assessment of precipitation trends in the
 upper Blue Nile River basin. *Stochastic environmental research and risk assessment*, 29,
 1751-1761.
- 520 TABARI, M. M. R. & TAVAKOLI, S. 2016. Effects of stepped spillway geometry on flow pattern 521 and energy dissipation. *Arabian Journal for Science and Engineering*, 41, 1215-1224.
- 522 TAMIRU, M. 2011. *Evaluation of Extreme Rainfall and Temperature Variability (In Upper Blue* 523 *Nile, Ethiopia)*. Addis Ababa University.
- TAPLEY, B. D., WATKINS, M. M., FLECHTNER, F., REIGBER, C., BETTADPUR, S., RODELL, M.,
 SASGEN, I., FAMIGLIETTI, J. S., LANDERER, F. W. & CHAMBERS, D. P. 2019. Contributions
 of GRACE to understanding climate change. *Nature climate change*, 9, 358-369.
- TOTARO, V., GIOIA, A. & IACOBELLIS, V. 2020. Numerical investigation on the power of
 parametric and nonparametric tests for trend detection in annual maximum series.
 Hydrology and Earth System Sciences, 24, 473-488.
- TURNER, J., MARSHALL, G. J., CLEM, K., COLWELL, S., PHILLIPS, T. & LU, H. 2020. Antarctic
 temperature variability and change from station data. *International Journal of Climatology*, 40, 2986-3007.
- VOGEL, R., ROSNER, A. & KIRSHEN, P. 2013. Brief Communication: Likelihood of societal
 preparedness for global change: trend detection. *Natural Hazards & Earth System Sciences*, 13.
- WANG, F., SHAO, W., YU, H., KAN, G., HE, X., ZHANG, D., REN, M. & WANG, G. 2020. Reevaluation of the Power of the Mann-Kendall Test for Detecting Monotonic Trends in
 Hydrometeorological Time Series. *Frontiers in Earth Science*, 8, 14.
- 539 WATSON, R. T., ZINYOWERA, M. C. & MOSS, R. H. 1996. *Climate change 1995. Impacts,* 540 *adaptations and mitigation of climate change: Scientific-technical analyses.*
- YAN, Y., XUE, B., YINGLAN, A., SUN, W. & ZHANG, H. 2020. Quantification of climate change and
 land cover/use transition impacts on runoff variations in the upper Hailar Basin, NE
 China. *Hydrology Research*.
- YUE, S., PILON, P. & CAVADIAS, G. 2002. Power of the Mann–Kendall and Spearman's rho tests
 for detecting monotonic trends in hydrological series. *Journal of hydrology*, 259, 254271.